MOISE: A PILOT EXPERIMENT TOWARDS LONG TERM SEA-FLOOR GEOPHYSICAL OBSERVATORIES

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Abstract

We describe the scientific purposes and experimental set-up of an international deployment of a 3 component broadband seismometer package on the ocean floor in Monterey Bay which took place during the summer of 1997. Highlights of this experiment were the installation, performed using a remotely operated vehicle (ROV), the underwater connection of the different components of the package, and the successful retrieval of 3 months of broadband seismic and auxiliary data. Examples of recordings of teleseisms and regional earthquakes are presented and the background noise characteristics are discussed, in comparison with those of near-by broadband land sites, current-meter data from the vicinity of the ocean bottom package, as well as pressure data from deeper ocean sites.

Introduction

The need for long term ocean floor observatories for the purpose of solid earth studies has now been widely recognized and several national and international efforts are underway to resolve the challenging technological issues associated with such deployments. In the summer of 1997, we conducted a pilot, target of opportunity international experiment in Monterey Bay (MOISE, Monterey Bay Ocean Bottom International Seismic Experiment, Stakes et al., 1997;1998), associating the resources and expertise of three teams: MBARI (Monterey Bay Aquarium Research Institute), a multi-institutional team from France (IPGP,DT/INSU: Division Technique de l'Institut National des Sciences de l'Univers, UBO), and the Seismological Laboratory at UC Berkeley. The primary goal of this experiment was to demonstrate the feasibility of deployment and intermediate term operation of several geophysical packages on the seafloor, using a remotely operated vehicle (ROV). The installation was successful, and 3 months of broadband seismic and various auxiliary data were acquired and processed. In what follows, we briefly describe the background motivation and technical aspects of the MOISE experiment and discuss the resulting broadband seismic data.

Scientific motivations

Due to the limited distribution of continents and ocean islands around the world, it has long been recognized that the coverage in land-based seismic and other geophysical observatories is not adequate to address many important scientific issues related to plate tectonics and the deep structure and dynamics of the earth. These have been described in detail in several documents (COSOD II, 1987; Purdy and Dziewonski 1988; Forsyth et al., 1995; Montagner and Lancelot, 1995) and we will only briefly summarize them here.

In spite of significant international efforts to increase the distribution of land-based state-of-the-art broadband seismic stations around the world (i.e. Romanowicz and Dziewonski, 1987), many areas of the southern hemisphere remain poorly sampled for global tomographic studies of the earth's deep mantle. This limits resolution in the distribution of lateral heterogeneity and anisotropy in the upper-mantle and transition zone, and therefore our ability to map the precise geographical distribution of dynamically critical upwellings and downwellings, and, for example, determine the precise nature of the boundary-layer at the 670 km discontinuity. Likewise, based on studies of specific patches sampled by currently existing data, the D" region has been recognized as the locus of strong lateral heterogeneity and anisotropy. Assessing whether these observations represent a singularity or a general property of D" and constraining their physical nature is critical to our understanding of the dynamics of this crucial boundary layer, yet we are limited by the global distribution of earthquake sources and receivers, even as the pool of continental stations has improved in the last ten years (figure 1). Better sampling of the inner core is also needed to improve our understanding of its structure and anisotropy and this necessitates installation of seismic stations at optimal locations with respect to seismic sources (for example, antipodal), and most of these are under water. The sparse distribution of seismic stations in the southern hemisphere also limits the azimuthal sampling of the source radiation in the case of southern earthquakes, and therefore limits our resolution of the corresponding source processes.

At the regional scale, California is in a particularly good position to benefit from the installation of off-shore permanent broadband seismic stations, to complement the land based network. Indeed, most of the existing stations are located on the east-side of the north-american/Pacific plate boundary and offer limited azimuthal coverage for moderate to large regional earthquakes on the main active faults (figure 2). They also provide limited location and source mechanism resolution for infrequent and generally smaller ($M < 4$) off-shore earthquakes on faults, such as the San Gregorio, that are known to have experienced significant (M6) events. Finally, off-shore broadband recordings of on-land earthquakes would provide much needed data to constrain the three-dimensional structure of the crust and thereby improve our understanding of the nature and evolution of this plate-boundary. Since large regional events are rare, deployment of ocean bottom stations for long periods of time (years) is necessary for this purpose.

There are several technical challenges that need to be addressed before permanent ocean bottom seismic observatories are deployed routinely: issues of optimum installation (borehole versus ocean floor), means of installation (manned or unmanned vehicles), sources of power supply and means of data retrieval are among the most critical ones.

One of the critical issues in the design of these ocean bottom observatories is how to optimize installation in order to minimize the background noise in the frequency bands most relevant to the recording of regional and distant earthquakes. Background noise on the sea-floor has been well documented in the frequency band of interest (10^{-3} to 10 Hz) from a combination of standard short-period OBS deployments and differential pressure-gauge measurements (e.g. Webb, 1997). Notably, such an experiment was conducted in San Francisco Bay in the late 1960's (Sutton and Barstow, 1990). However, measurements from 3 component broadband seismometers have been few. The main problem is the challenge of installation of broadband seismometers on the ocean floor, since they require precise leveling and centering that is difficult to achieve remotely. In recent years, two experiments successfully retrieved broadband seismic data acquired from modern seismometers in the atlantic ocean (Montagner et al., 1994b) and in the Japan Sea (Suyehiro et al., 1992). Due to technical difficulties, in both experiments, only several days of ocean bottom data were retrieved. The issue of whether installation directly on the ocean floor (with partial or complete burial in the sediments) or in boreholes is preferable has not been completely resolved yet and is one of the objects of the ongoing OSN-1 experiment (Orcutt and Stephen, 1993; Purdy, 1995).

MOISE is therefore one among several pilot experiments designed in the spirit advocated by the International Ocean Network program (ION, Suyehiro et al., 1995) as well as the U.S. Ocean Seismic Network program (OSN, Purdy and Dziewonski, 1995), and aimed at defining the technology and optimal deployment procedures appropriate for the future installation of permanent ocean bottom geophysical observatories. In spite of high noise levels encountered, the MOISE experiment successfully demonstrated the feasibility of an intermediate term stand-alone deployment on the ocean floor.

Technical Description of MOISE experiment

The site of the MBARI experiment was chosen in February 1997, on a flat plateau area called Smooth Ridge in Monterey Bay, just beyond the active San Gregorio fault (figure 3), 40 km off-shore, at a water depth of 1015m. At the time of the site selection, the ocean bottom currents were very low, indicating that this location should be optimally "quiet" from the seismic point of view. The deployment of the scientific packages was performed using the MBARI ship "Point Lobos" and the MBARI ROV "Ventana". The location of the experiment was accessible within 2 hours transit time from Moss Landing (CA), the MBARI headquarters.

The central instrument package of MOISE was a 3 component broadband seismic package composed of Guralp CMG-3 sensors mounted on leveling gymbals. The electronics were adapted for this ocean floor deployment by DT/INSU. The seismometers had been previously deployed in the mid-atlantic ocean, during the OFM experiment (Observatoire Fond de Mer, Montagner et al., 1994a,b). The seismic package includes a 16-bit, gain-ranged digitizing system, an automatic re-centering system as well as a self-leveling system driven by a small cpu. The cylindrical aluminium housing was designed by the DT/INSU and outfitted with handles for manipulation by the ROV and the male side of an 8-pin Nautile connector (figure 4). During the deployment, the seismic package was connected underwater to a lithium battery package as well as to the recording system. The recording system was designed and built by Scripps Institution of Oceanography (SIO) for ocean bottom applications ("El Cheapo", J. Orcutt, personal communication). The "El Cheapo" software was modified for the MOISE experiment so that it would allow simultaneous connections to the seismic sensor package and the ship. In this way, it was possible to establish communication with the sensor package directly from the ship, which was done three times: at installation and retrieval time, and, also, one month into the duration of the experiment, to verify the good functioning of the seismometers and data recording. During this intermediate visit to the site, the ROV established a connection from the ship to the "El Cheapo" recorder, and it was possible to retrieve samples of data, as well as connect a lap-top computer aboard the ship to the seismometer system, in order to relevel and recenter the seismometers. The software also allowed to instruct the seismometer, at the beginning of the experiment, to unlock the gymbals once installed, and, at the end of the experiment, to shut down and lock the gymbals.

In order to achieve good coupling of the seismometer system with the ground, efforts were made to partially bury it in the sediments. Site preparation included placing a large weight to compact the sediment several weeks before deployment. At the time of deployment, a PVC cylinder was first brought to the site and 75% buried in the soft sediment, which was removed from inside the cylinder using a shovel manipulated by the ROV. The seismometer package was then lowered into this hole by elevator. The gap left between the seismic package and the PVC cylinder was filled with 3mm glass beads to consolidate the installation and allow easy recovery at the end of the experiment. The seismic package had a rather elongated shape (Figure 4), partly because the underwater connectors were mounted on its top and, once installed, was protruding about 70cm above the sea-floor, thus presenting an unfortunately large surface to the bottom ocean currents.

Difficulties during installation due to unexpected levels of bottom water currents and the hovering position of the rover during the installation of the heavy (70 kg in water) seismic package, prevented the Ventana operators from positioning the package completely vertically and the resulting tilt of the seismometers was near the limit of linearity for the instrument specifications. It also prevented them from installing an aluminium "hat" over the package, which had been designed by DT/INSU to shield it from direct action by currents. The most suspenseful part of the installation was the fitting of the underwater connectors between the seismometer and the recording system, accomplished while the Ventana was hovering in strong (up to 20cm/s) currents. It took 2.5 hours to establish the first acrobatic connection. The following ones were performed faster, as the Ventana operators gained experience and the dataflow was successfully established, via RS232 connections, first to the ship, and then to the recording system, the latter for the 3 months duration of the experiment.

In addition, a CTD/pressure gauge (on loan from Curt Collins), with a support frame and anchor, as well as an S4 current meter, were deployed on a short mooring in the vicinity of the seismic package for the entire duration of the experiment. These were stand-alone instruments In particular, the current meter data proved invaluable in understanding the sources of seismic background noise during this deployment. A self- contained electromagnetic package containing magnetometers, electrometers and an internal data logger designed at UBO (Montagner et al., 1997) was deployed at the same site during the same time period. The resulting data will be described in a separate publication. Finally, several other instruments were deployed in Monterey bay during the same time, as part of a separate experiment (MBARI Margin seismology project): 3 standard three channel 4.5 Hz OBS's and a single channel hydrophone, to monitor microseismic activity of the San Gregorio and Monterey Bay Fault zones.

Seismic Data Analysis and results

Three component broadband seismic data were acquired continuously on the ocean floor from 06/21/97 through 09/11/97, at a sampling rate of 20 samples/sec. Several regional earthquakes of magnitude 3.5 and larger as well as several large teleseisms were well recorded during that time period. Comparison of the records obtained at the ocean-bottom site (MOIS) with those of near-by land sites of the Berkeley Digital Seismic Network (BDSN) provide useful insight into the quality of the data. Figures 5 and 6 show examples of recordings at MOIS of a regional and a teleseismic event respectively, compared to those of BDSN station FARB, a noisy island site (Figure 3), as well as SAO, the closest BDSN continental site. The regional event data are raw, but the teleseismic ones were band-pass filtered between 5 and 50 sec, a frequency band of minimum noise (or low-noise "notch") as documented from previous ocean-bottom experiments (e. g. Webb et al.,1991). Background noise levels were found to fluctuate considerably in this period band and, not surprisingly, the best recordings were obtained during the quietest periods, when noise in the minimum noise window was comparable to that observed commonly at nearby land sites. In figure 6, we note that the P wave and its multiples as well as the S wave and Rayleigh waves are well recorded on the vertical component of MOIS. The S wave and Rayleigh waves are visible on the east component bu the north component is particularly noisy during this time interval.

Figure 7 presents a comparison of noise spectra recorded at MOIS and the land stations JRSC and SAO at different times during the same day (09/01/97). At periods shorter than 10 sec, the noise is stable for this time period at all three stations. The microseismic peak is somewhat wider at MOIS than at SAO, however, noise decreases rapidly at higher frequencies, and, at 1 Hz, reaches levels comparable to those at SAO and lower than those at JRSC. At periods longer than 5 sec, the "low-noise notch" is particularly prominent on the vertical component of MOIS. Beyond 30-50 sec, and in contrast to the land stations, noise at MOIS increases rapidly on all three components. Noise fluctuations at low frequency are large and can exceed 30dB on the vertical and horizontal components alike. In the case of land stations, noise fluctuations at low frequencies have been documented to be related to atmospheric pressure fluctuations, particularly on the horizontal components. On the ocean floor, as documented from pressure measurements, noise in the "low-noise notch" is known to fluctuate with bottom current velocity (e.g. Webb et al., 1991). Investigation of the source of the large noise fluctuations at station MOIS in the low-noise window was made possible owing to the simultaneous recording of auxiliary data. In figure 8, we present the fluctuations of current velocity and seismic noise over a period of 24 hours, on a seismically quiet day (09/01/97). The variation with time of the theoretical solid tide at the MOIS location is also given for reference. As figure 8 illustrates, in the period band 10-50 sec, background seismic noise is often strongly correlated with the bottom current velocity, which exhibits large fluctuations that can in general be related to tides. Figure 9 shows a similar plot on a day when the correlation was poorer, due primarily to the occurrence of a teleseism (Western Australia, Ms 6.0; origin time 09:20 UTC). This increased the levels of power spectral density in the seismic signal during a time whe the current velocity happened to be low.

The current strength turned out to be unexpectedly high during the summer at the MOISE site, a situation which prevailed for the entire duration of the experiment and which has been attributed, at least partly, to the presence of the El Nino phenomenon. The strong influence of bottom currents on the background seismic noise is easily explained by the experimental set-up, with the tilted seismic package protruding significantly above the sea-floor. On the other hand, during quiet periods, the background noise between 10-30 sec period -a period appropriate for observations of teleseismic long period body waves and regional surface waves - appears to be comparable to that observed at nearby continental sites. This indicates to us that, at least in the long period band, proper installation in the relatively soft ocean- floor sediment, with the instrument completely buried and shielded from currents would likely result in a site quality comparable to that achieved on average on land sites (e.g. Sutton and Duennebier, 1988; Duennebier and Sutton, 1995). This conclusion is in agreement with that reached during the OFM experiment, in which the ocean-bottom installed broadband seismic package was not completely buried in the ground, however, it was protected from currents by an aluminium shield similar to that initially planned for the MOISE experiment (Beauduin et al., 1996).

Figure 10 compares noise levels on the vertical component at MOIS and SAO for three different epochs during the experiment. In addition to fluctuations at periods longer than the microseismic peak, as previously described, we note variations in the noise in the microseismic band that appear to be correlated between the two stations, with higher levels of noise on the ocean floor. Noise in this period band has been documented to be strongly correlated with weather patterns (e.g. Webb, 1997) which should affect the ocean bottom station and near by land stations in similar ways. At the shortest periods analyzed (1-5Hz) noise at MOIS appears to be more stable than at SAO, which explains the high quality of the short-period filtered MOIS records obtained for the regional event shown in Figure 5. At periods shorter than 1Hz, there appears to be again periods of high correlation of the MOIS background noise with the bottom current, in agreement with previous results from pressure gauge deployments (Webb, 1991). Figures 11a and b show spectrograms of the correlation of power spectral density (vertical and north component respectively) with the bottom current velocity . During many days, the correlation on both sides of the microseismic peak reached values in excess of 0.9. Unfortunately, it was not possible to take advantage of this observation and subtract the current-correlated signal from the seismic data in the long period band of interest for earthquake observations (10-50 sec), because the sampling rate of 300 sec of the current-meter data was insufficient. The correlation with current velocity notably persits at low frequency beyond the "low-noise notch". Infragravity waves are thought to be the main cause of rising levels of noise at periods larger than 50sec in deep ocean pressure measurements (Webb, 1991). During the intervals of time when the correlation of MOIS noise with bottom current is not strong, such infragravity waves may be a dominant cause of noise down to periods of 30 sec, given the relatively shallow location of the seismic package. This could explain the relatively narrow "low noise notch" in the MOISE data, as compared to data from differential pressure gauges in the deep ocean (e.g. Beauduin et al., 1996). We note from figure 7 that the daily noise fluctuations at periods longer than 10 sec present some coherency between the MOIS site and the nearby land sites, indicative of a complex regional interaction of tides, currents and atmospheric disturbances.

Conclusions

Through the acquisition of 3 months of continuous three component broadband seismic data in Monterey Bay, the MOISE experiment has demonstrated the feasibility of deployment of long term autonomous broadband seismic packages on the ocean floor using ROV technology and the usefulness of auxiliary instrumentation, particulary current-meters, to understand the sources of background noise. It was indeed found that, apart from the microseismic peak frequency band, the background noise on all three components was highly correlated with current velocity. At frequencies higher than 1Hz, noise levels were comparable to or better than those recorded on land stations. The microseismic peak level was 20dB higher than on nearby land stations, comparable on all three components. The low noise "notch" at periods between 10-30 sec was particularly marked on the vertical component, where, during quiet periods, noise levels were comparable to those at nearby SAO land station, allowing good intermediate frequency recordings of several regional and teleseismic events. These results indicate that proper installation with seismometer package entirely buried in the sediment should have provided data of quality, comparable to that achieved at near-coastal land stations, on both sides of the micro- seismic peak.

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FIGURE CAPTIONS

Figure 1. Typical coverage of the D" region currently achievable with data from global broadband networks. Lines indicate sampling of D" by S-diffracted waves and dots indicate locations of reflection points for ScS, ScS2 and ScS3 phases achieved in the global tomographic model SAW12D (Li and Romanowicz, 1996). Note the much poorer coverage in the southern hemisphere.

Figure 2. Current distribution of land-based broadband seismic stations in the western US, and tentative distribution of permanent broadband ocean floor observatories of the future that would complement the land-based network for a more balanced coverage of the western US plate boundaries.

Figure 3. Location of the MOIS deployment, 40 km off-shore in Monterey Bay, shown relative to the closest land broadband stations of the Berkeley Digital Seismic Network (BDSN). Station FARB is an island site equipped with a set Guralp CMG40T broadband seismometers. Other stations operate Streckeisen STS-1's or STS-2's. Station BKS is located near our central processing site at UC Berkeley in the San Francisco Bay Area. Major faults are indicated by thick lines. Broken lines indicate approximate traces of major off-shore faults, such as the San Gregorio fault just East of the MOIS site.

Figure 4. Photograph of the seismometer package before installation. At the top of the structure is the underwater connector linking the seismometers to the recording system located in its vicinity.

Figure 5. Example of raw broadband data for a regional earthquake (Nevada, distance $\tilde{4}00$ km) recorded at MOIS, a nearby island station, FARB and and a nearby continental station, SAO.

Figure 6. Example of data for a teleseism for the same stations as in Figure 5. Note the clear recording of body and surface waves on the vertical component of MOIS. Data were

filtered between periods of 5 and 60 sec.

Figure 7. Comparison of noise spectra at land stations SAO and JRSC and ocean bottom station MOIS, computed using 1 hour of data successively starting at 12:35 (solid line), 16:35 (broken line), 20:35 (dotted line) UTC on a day free of significant earthquakes (08/18/97). Columns 1-3 show noise on the East, North and vertical component respectively.

Figure 8. Fluctuations as a function of time of day (on 09/01/97) of bottom current velocity (black) and direction (arrows, north is towards top of figure) compared to those of MOIS vertical component noise power spectral density in the period ranges 10-20 sec (red), 20-30 sec (blue) 30-50 sec (green). Power spectral density (PSD) scale is on the left and velocity scale on the right.

Figure 9. Same as Figure 8 on August 10, 1998. The increase in power spectral density around 10:30 am UTC coincides with the arrival of surface waves from an Ms 6.0 earthquake in western Australia.

Figure 10. Comparison of PSD of seismic noise at SAO and MOIS (vertical component) at 3 epochs during the experiment: July 11 (solid line), August 18 (broken line), and September 1 (dotted line), 1997.

Figure 11. Correlation of current velocity and power-spectral-density of vertical (a) and north (b) component at MOIS site as a function of time and period. High correlation is observed on both sides of the microseismic peak on both vertical and horizontal components. Low correlation can sometimes, although not systematically, be associated with times of occurrence of large regional (M>4.5) or teleseismic (Ms>6.0) earthquakes. Such is the case, for example on day 207.