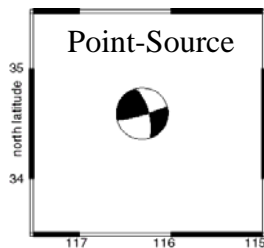
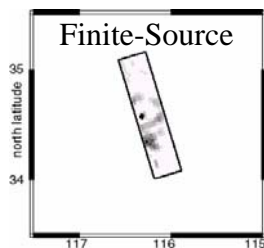


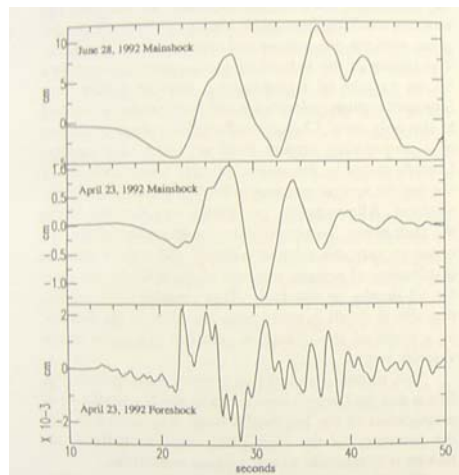
# Earthquake Kinematics & Dynamics

- Kinematics
  - Study of timing of seismic moment release
  - $M(t) = \mu A(t)u(t)$ 
    - Rupture velocity
    - Slip velocity / rise time
- Dynamics
  - Study of physical mechanisms governing earthquake rupture and seismic moment release
    - Initial stress state
    - Yield stress
    - Static & dynamic coefficient of friction
    - Frictional constitutive law
    - Pore pressure effects & melting

## Sources



Magnitude



## Kinematic Description of Finite Source Process

- Spatio-Temporal Descriptions of:
  - slip, rise time (slip velocity), and rupture speed
- Haskell Source Model
- Representation Theorem Models

## First Point-Source Kinematics

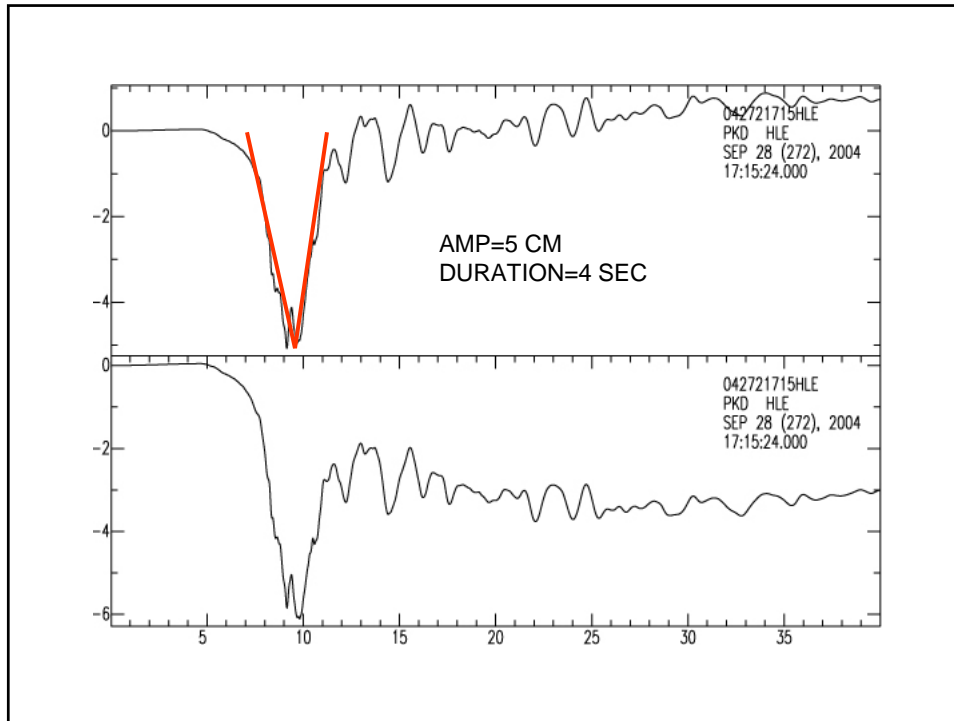
$$u_i(r,t) = \frac{2 \cdot R_i \cdot \dot{M}(t - r/\beta)}{4\pi\rho r\beta^3}$$

$$\int u_i(r,t) dt = \frac{1}{2} A \tau = \frac{2 \cdot R_i \cdot M_0}{4\pi\rho r\beta^3}$$

$$M_0 = \frac{(\pi\rho r\beta^3) A \tau}{R} = 3.14 * 2.67 * 8 \cdot 10^5 * (3.5 \cdot 10^5)^3 * 5 * 4$$

$$= 5.7 \cdot 10^{24} \text{ dyne cm}$$

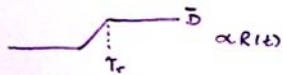
$$M_w = 5.8$$



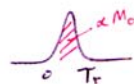
## What is a Reasonable form for the Moment Rate Function?

Types of slip time functions

$D(t)$



$\dot{D}(t)$

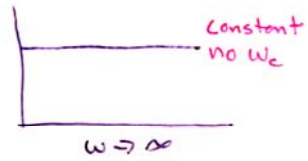


# Time & Frequency Domain Equivalents

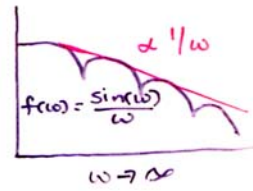
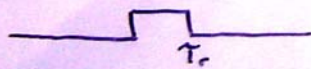
$M(t)$   
Delta Function



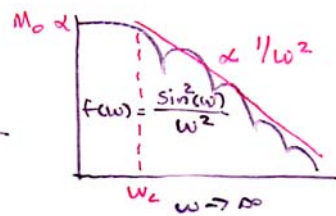
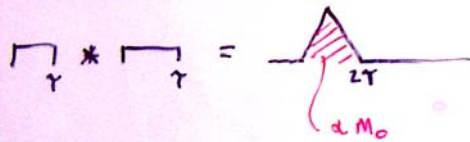
$|M(\omega)|$



Boxcar

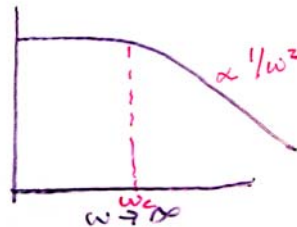


convolution of two Boxcars

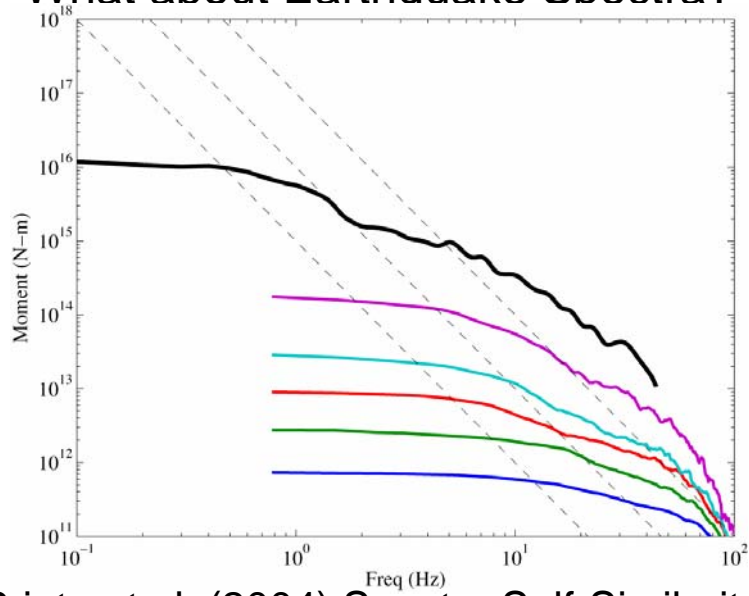


Omega squared model

$m(t) \propto t e^{-t/\tau}$

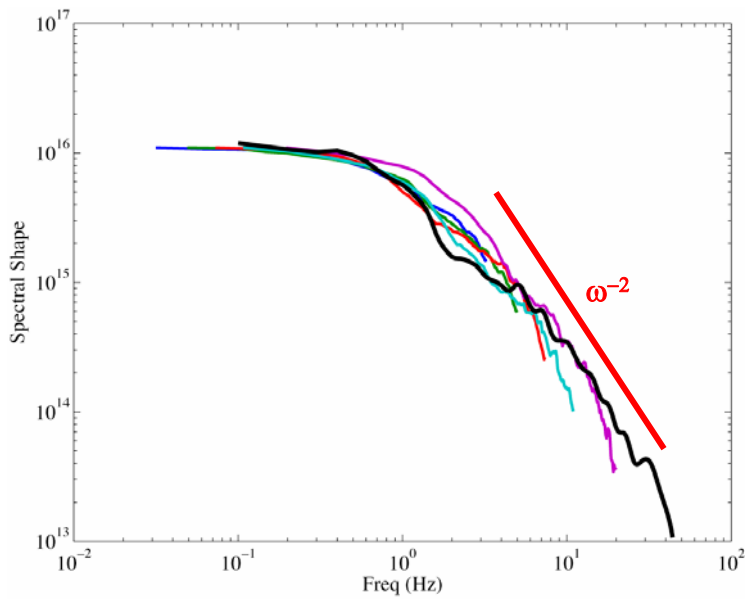


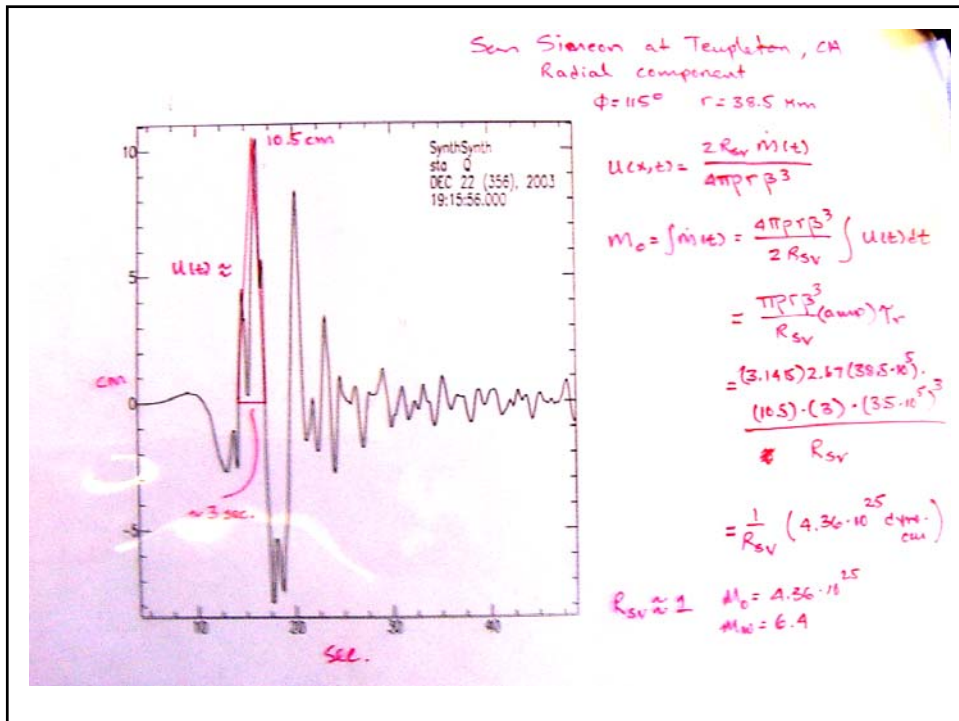
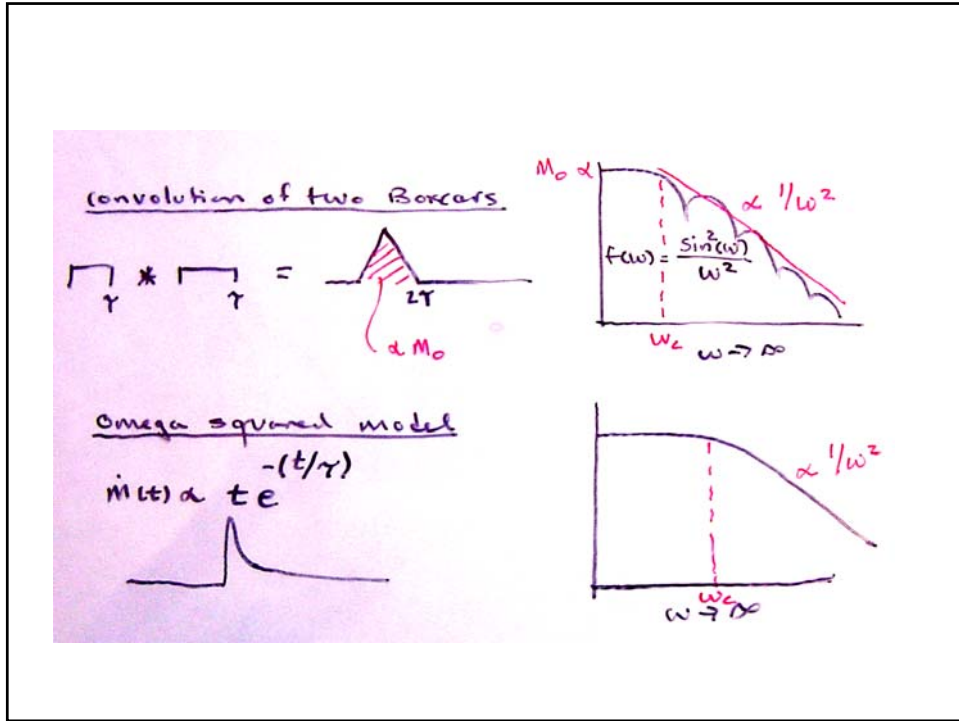
## What about Earthquake Spectra?

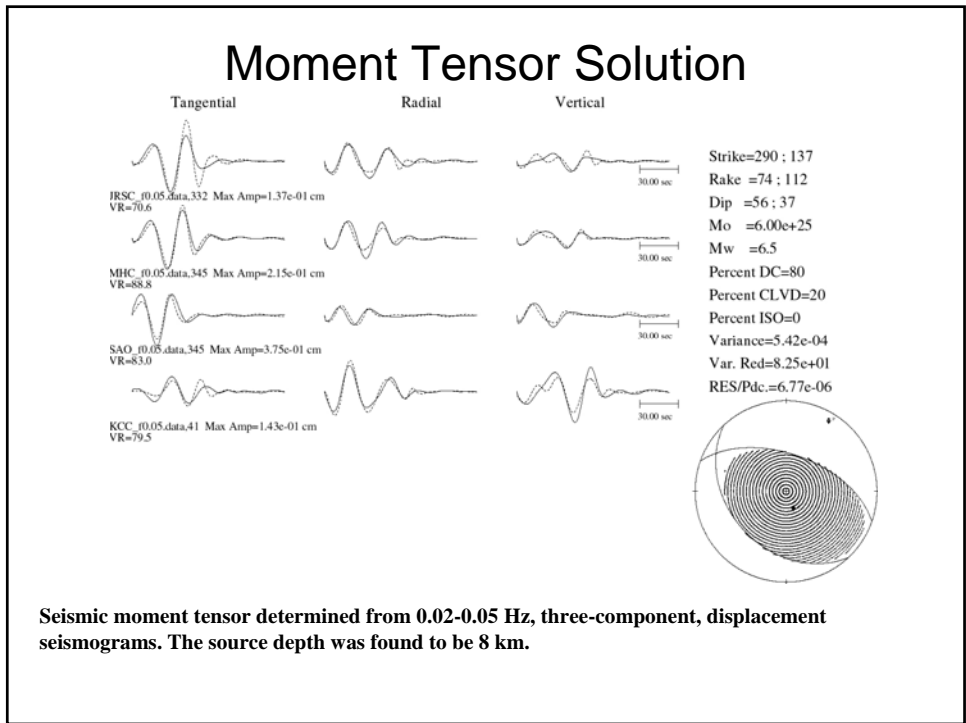
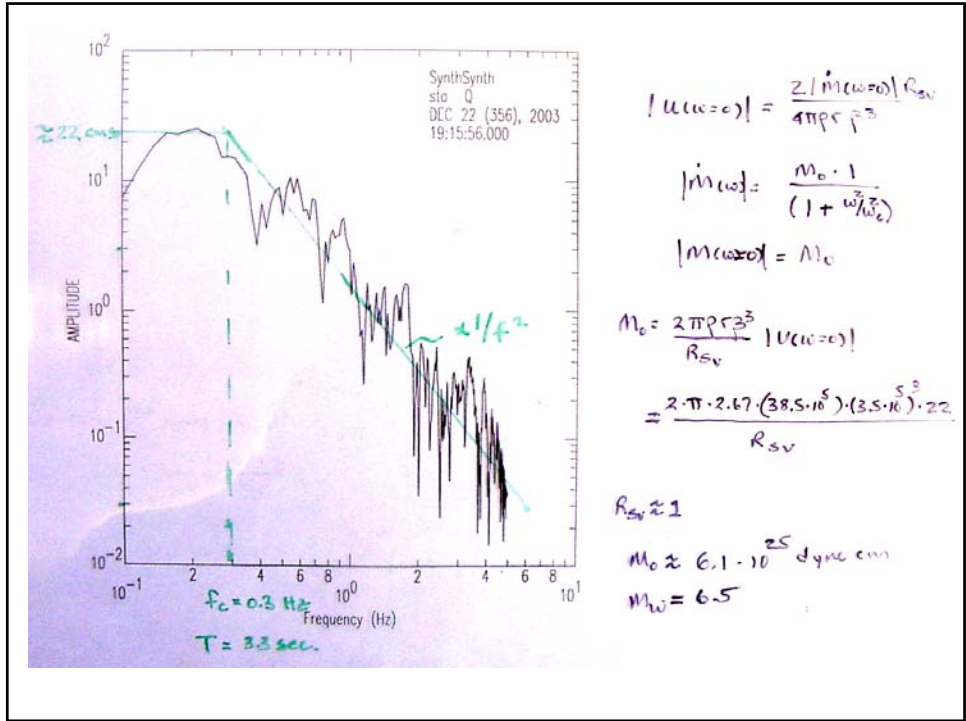


Prieto et al. (2004) Spectra Self-Similarity

## Prieto et al. (2004) Spectra Self-Similarity







Haskell Source Model assumes

- constant dislocation rise time
- uniform fault growth

Start with the far-field solution

$$u(r, t) = \frac{2R_{SH}}{4\pi\rho\beta^3r} \dot{M}(t-r/\beta)$$

$$M(t) = \begin{cases} 0 & t < r/\beta \\ \rightarrow M_0 & t \rightarrow \infty \end{cases}$$

$$M_0 = \int_0^{\infty} \dot{M}(t) dt$$

First break  $M(t)$  into 2. terms that depend on time.

1) slip rate  $\dot{D}(t)$  where  $\bar{D} = \int_0^{\infty} \dot{D}(t) dt$

2 Fault Length  $L(t) = \int_0^L \delta(t-x/v_r) dx$

then  $\dot{M}(t) = \mu W L(t) * \dot{D}(t)$

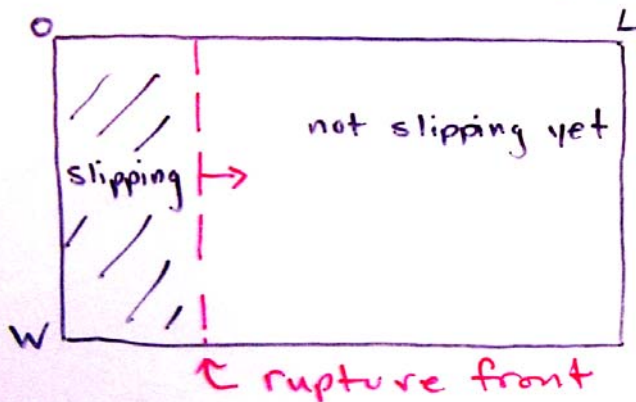
with a change of variable  $z = t - x/v_r$

$$dz = -\frac{1}{v_r} dx$$

the  $L$  integral simplifies



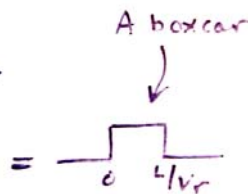
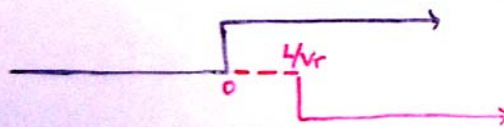
and a "ribbon" rupture model



the  $L$  integral simplifies

$$\dot{M}(t) = \mu W \dot{D}(t) * \int_t^{t-L/v_r} \delta(z) (-v_r) dz$$

$$= \mu W \dot{D}(t) * (v_r H(z)) \Big|_{t-L/v_r}^t$$



If it is assumed that  $D(t)$  behaves like a ramp function where total offset is achieved at time  $\tau_r$  then  $\dot{D}(t)$  is a boxcar with duration  $\tau_r$ .

So the Haskell solution is the convolution of two boxcars,  $B(\tau_r)$  and  $B(L/v_r)$

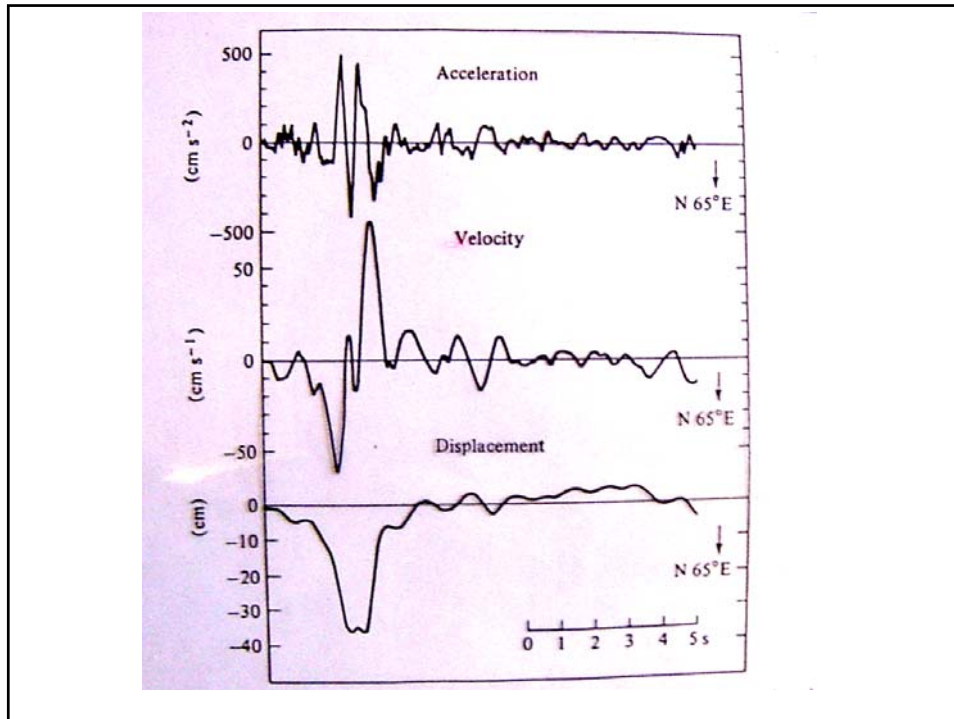
So the Haskell solution is the convolution of two boxcars,  $B(\tau_r)$  and  $B(L/v_r)$

If  $\tau_r = \tau_d$  the convolution yields a triangle

If  $\tau_r \neq \tau_d$  the convolution yields a trapezoid (see Box 9.2)



How reasonable can this be?



## Directivity

Difference in arrival time from opposite ends of a fault

$$t_{abc} - t_{ac} = \tau_d$$

$$\tau_d = \frac{L}{v_r} + \frac{r - L \cos \theta}{c} - \frac{r}{c}$$

$$\tau_d = \frac{L}{v_r} - \frac{L \cos \theta}{c}$$

$c$  is  $\alpha$  or  $\beta$  for

$P$  or  $S$  waves

$$v_r \approx 90\% \beta$$

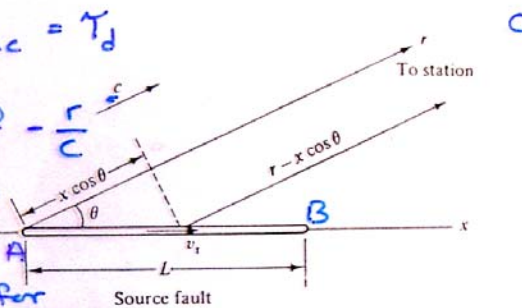
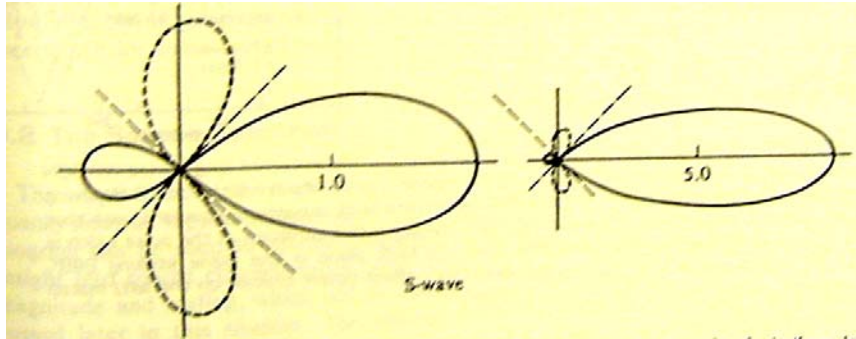


FIGURE 9.8 Geometry of a rupturing fault and the path to a remote recording station. (From Kasahara, 1981.)

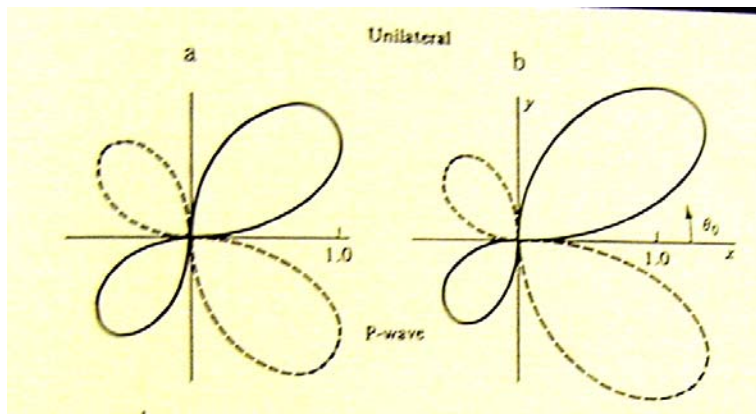
## Directivity & Radiation Patterns



$Vr/\beta = 0.5$

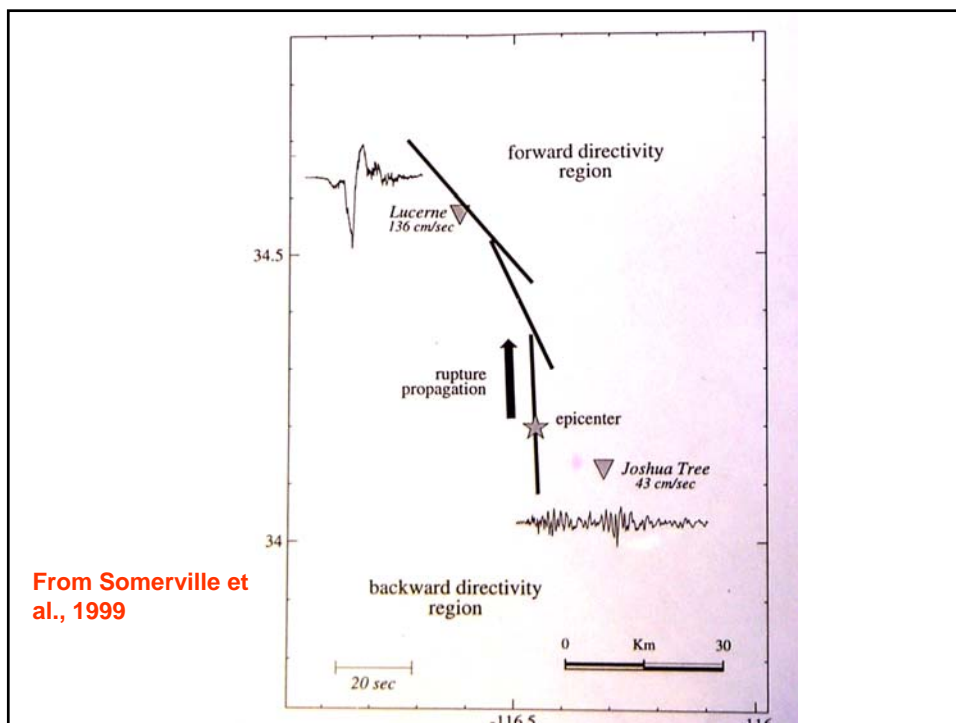
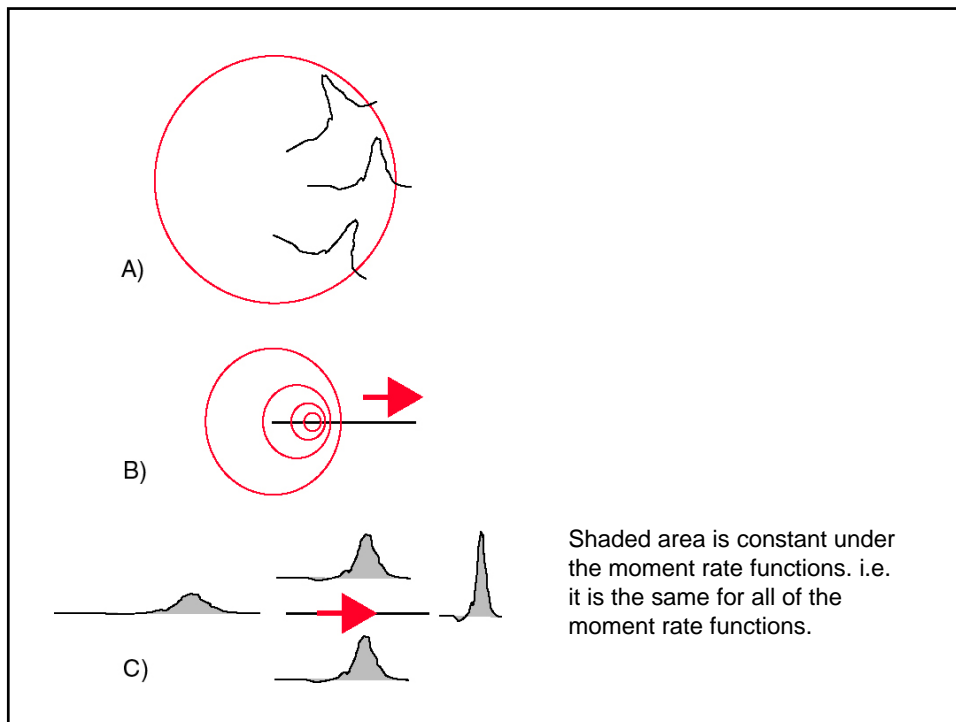
$Vr/\beta = 0.9$

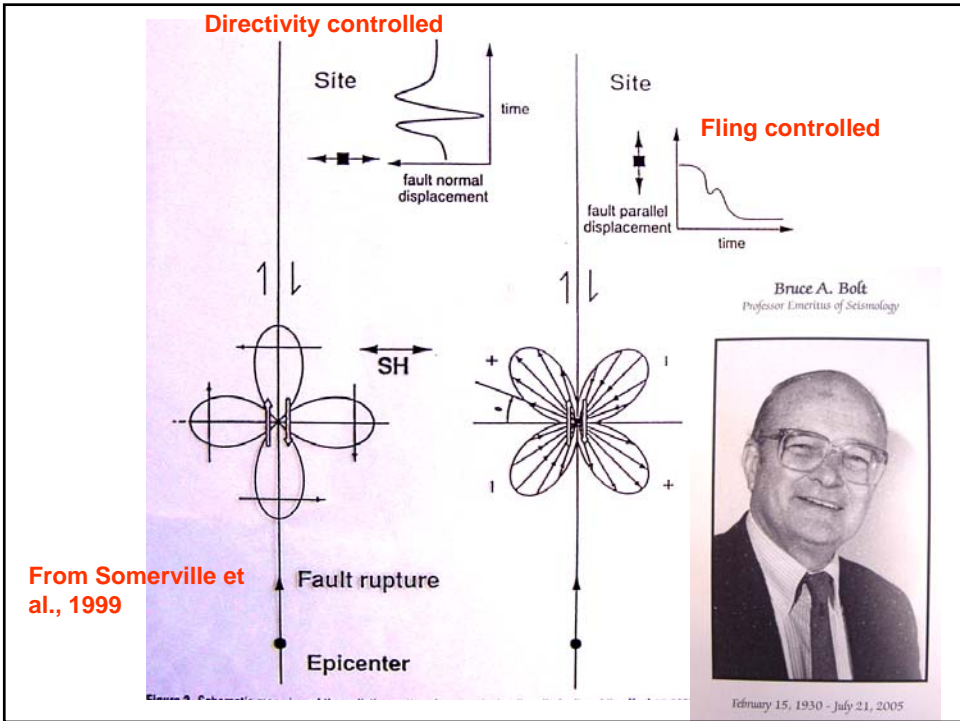
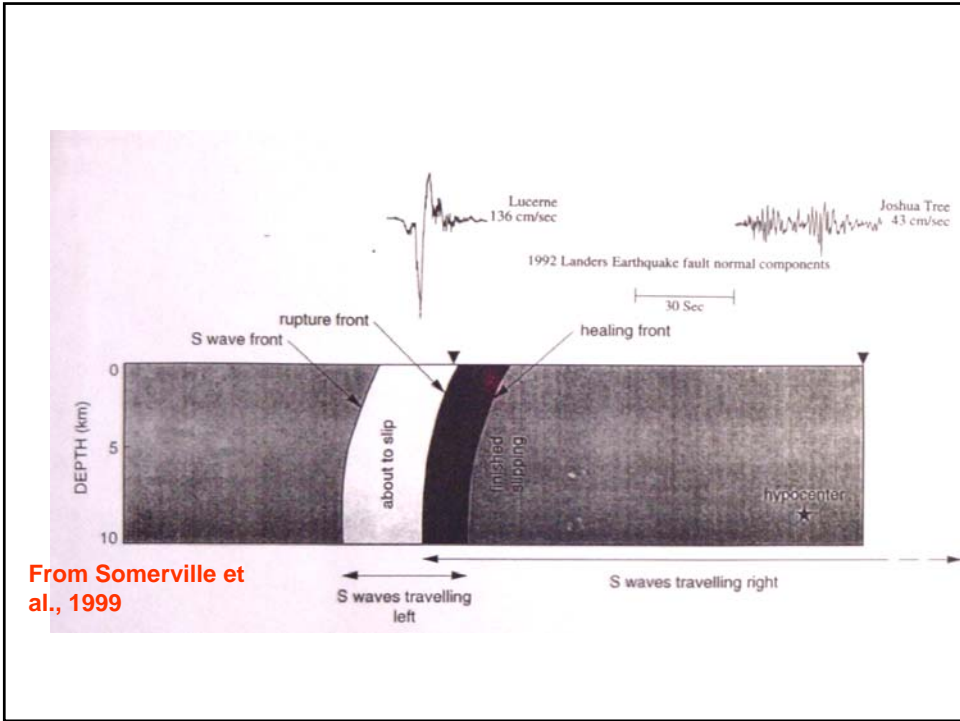
## Directivity & Radiation Patterns



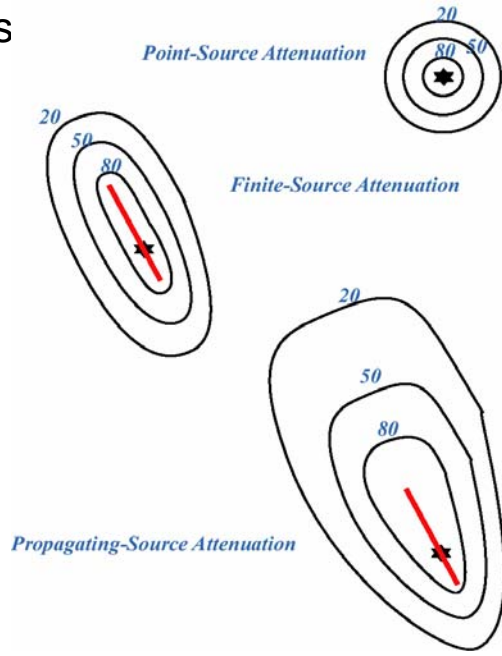
$Vr/\beta = 0.5$

$Vr/\beta = 0.9$



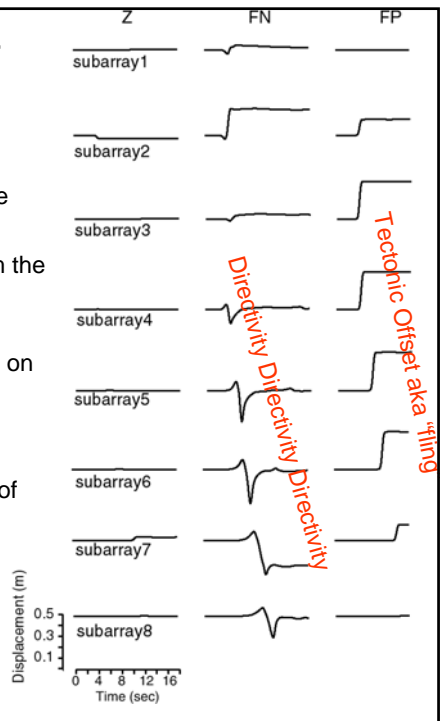


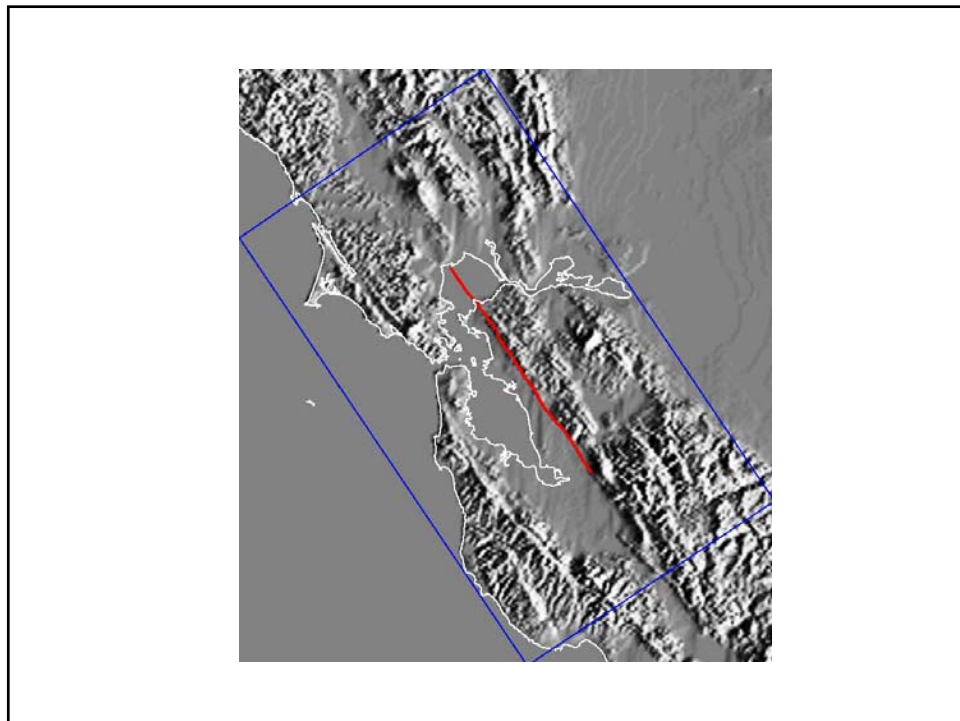
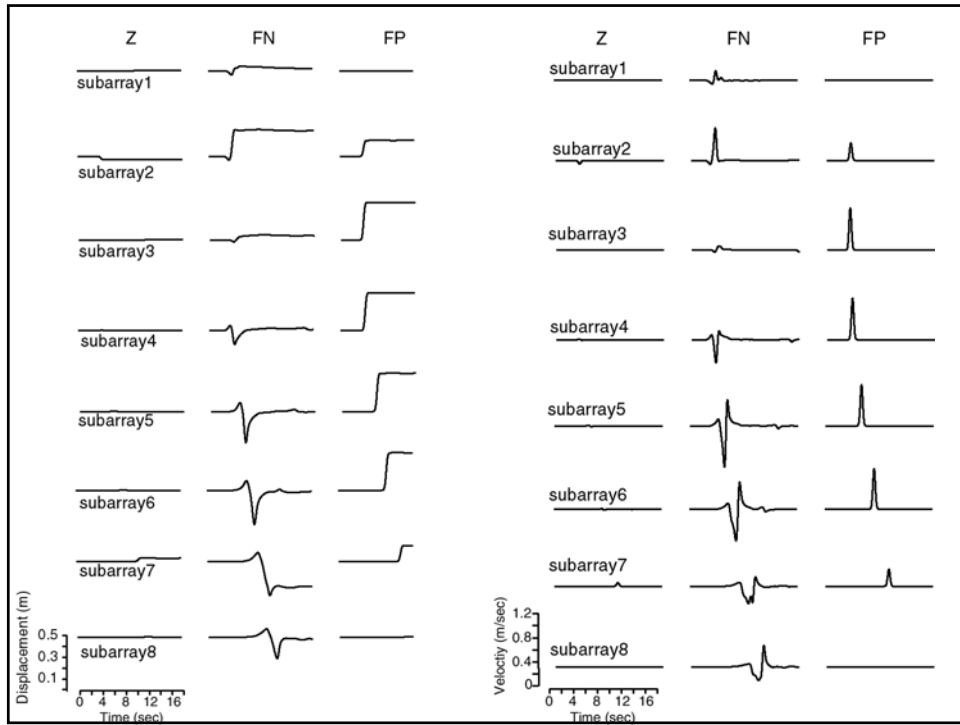
# Attenuation Models



## Near-Fault Ground Motions – Caltrans Project

- Unilateral, uniform slip, strike-slip rupture
- Computed using finite-differences
- Recording stations are located 15m from the fault
- Directivity builds displacement response on the FN component.
- The displacement (fling) on the FP component is constant along the length of the fault

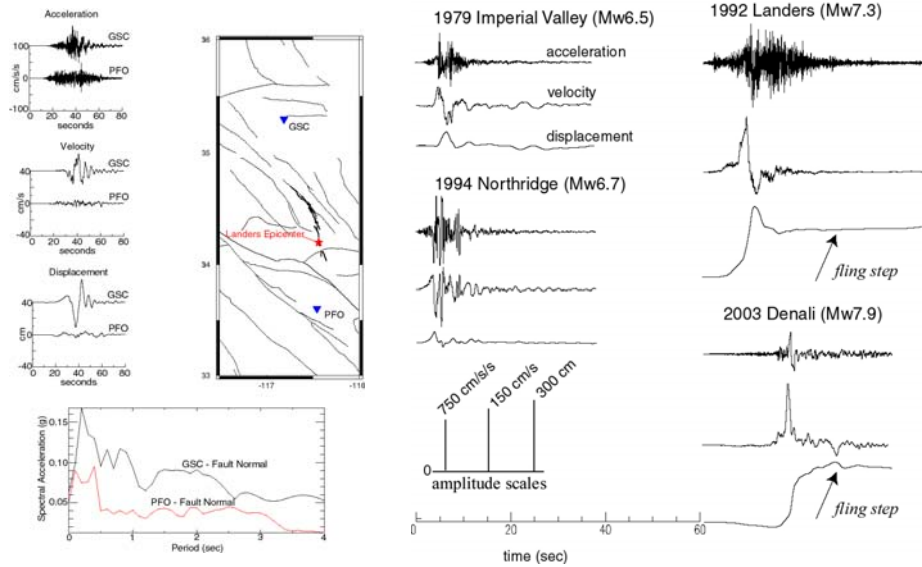


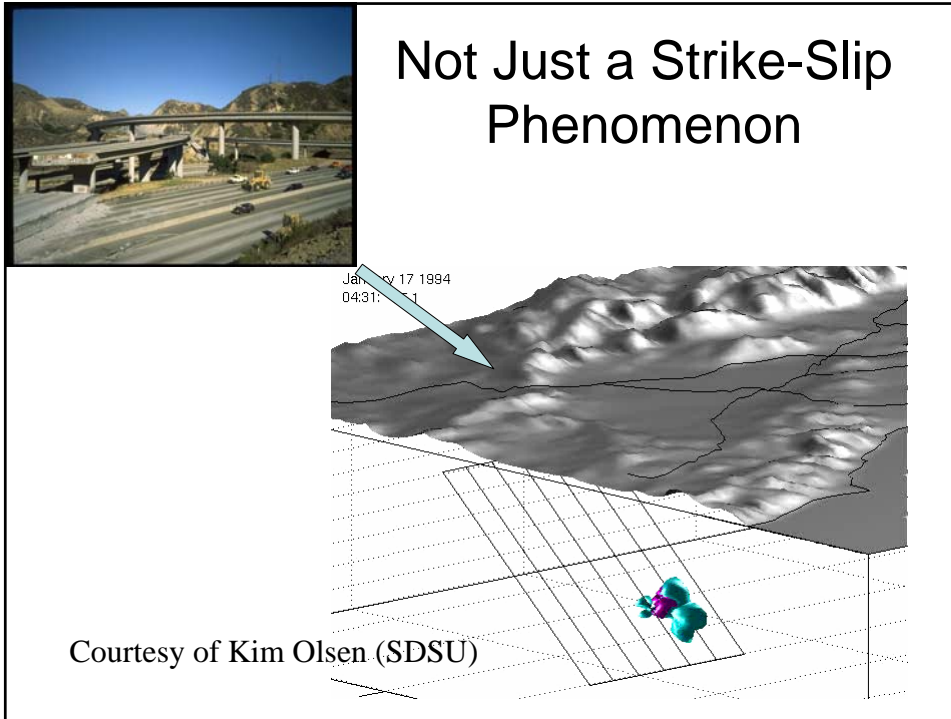




1906 Rupture at Tomalas Bay

Is Directivity & Fling Observed?





## Some Examples of Finite-Source Inversions

$$u_n(t, \vec{x}) = \int d\tau \iint [u_i(\vec{\zeta}, \tau) \hat{v}_j C_{ijkl}] \cdot G_{nk,l}(\vec{x}, t - \tau; \vec{\zeta}, 0) d\Sigma$$

$$u_n(t, \vec{x}) = \int [u_i(\tau) \hat{v}_j C_{ijkl} \Sigma] \cdot G_{nk,l}(\vec{x}, t - \tau) d\tau$$

Spatial point-source

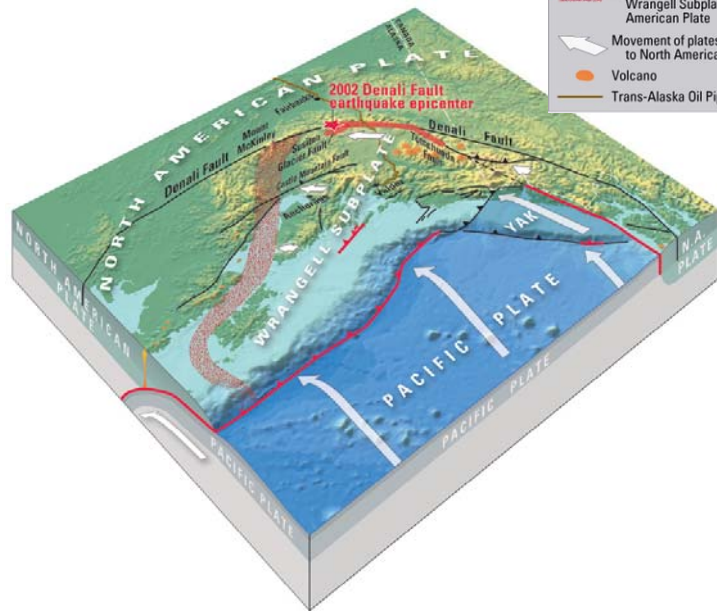
$$u_n(t, \vec{x}) = [u_i \hat{v}_j C_{ijkl} \Sigma] \cdot G_{nk,l}(\vec{x}, t)$$

Spatial and temporal point-source

$$u_n(t, \vec{x}) = M_{ij} \cdot G_{ni,j}(\vec{x}, t)$$

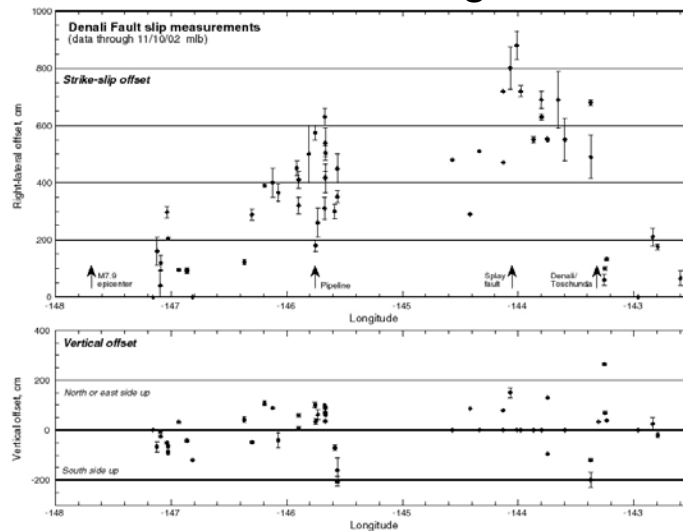
M has units of moment. i and j refer to directions of forces and derivatives. i.e. they define couples

# Tectonic Setting





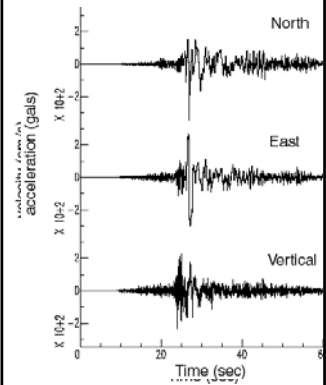
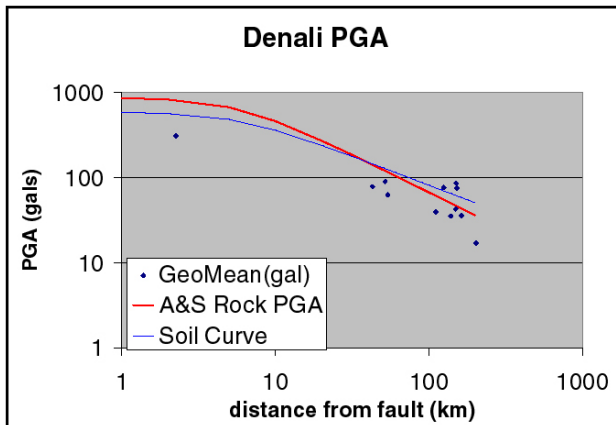
## Surface Faulting Data



Surface faulting data was provided courtesy of the Denali Earthquake Geologic Field Team. The surface offsets indicate that the rupture process was principally right-lateral strike-slip, although some vertical offsets were observed.



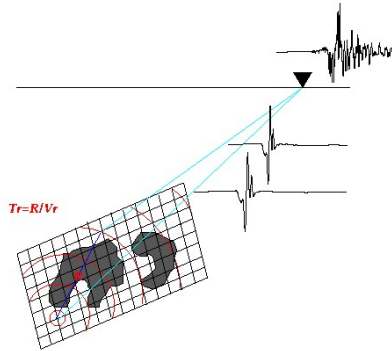
## Strong Ground Motions



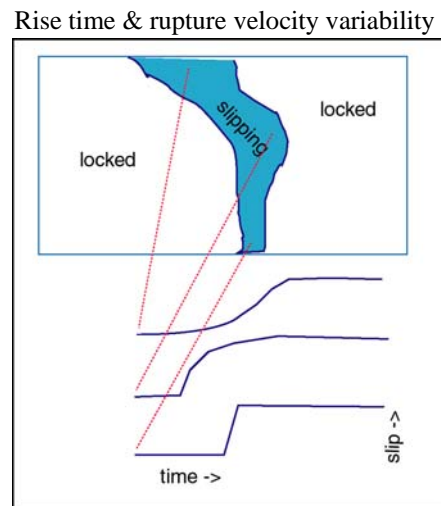
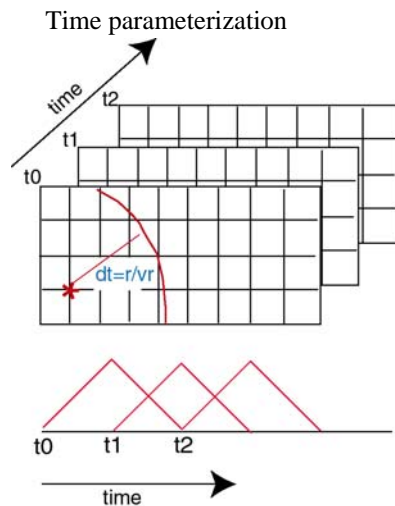
## Kinematic Method

$$U_n(x,t) = \int d\tau \iint [u_i(\xi,t) \hat{n}_j c_{ijpq}] \cdot G_{np,q}(x,0;\xi,t-\tau) d\Sigma$$

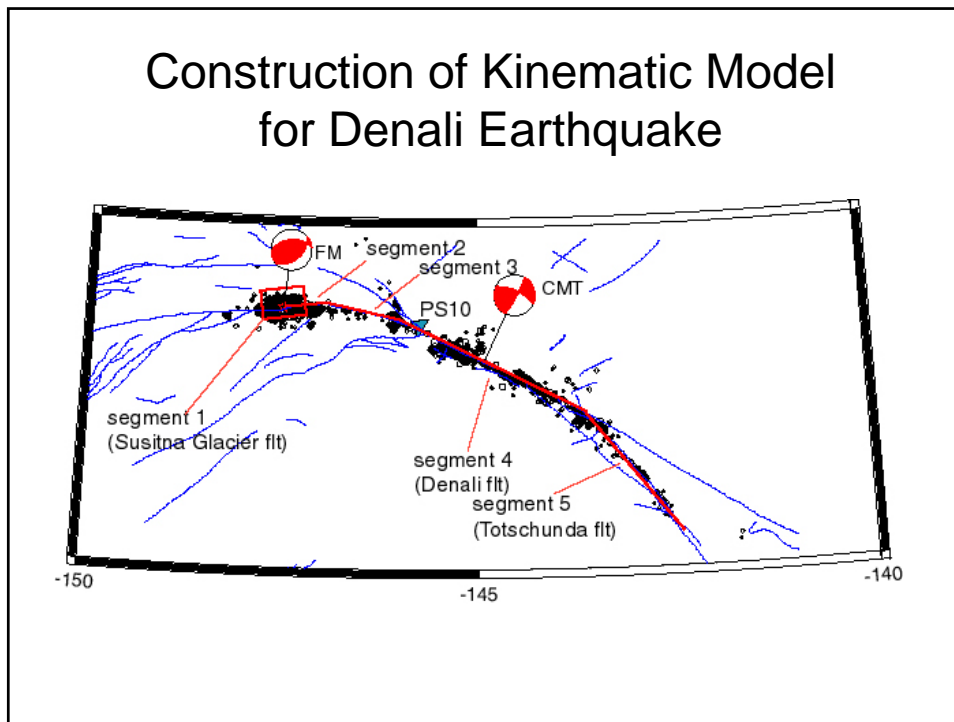
- Extension of Hartzell and Heaton (1983)
  - Use broadband displacement waveform data (0.01 to 5.0 Hz)
  - Use horizontal GPS deformation data
  - Apply surface slip constraining equation
  - Apply slip positivity, seismic moment minimization, and smoothing constraining equations
- Solve for spatial and temporal distribution of fault slip



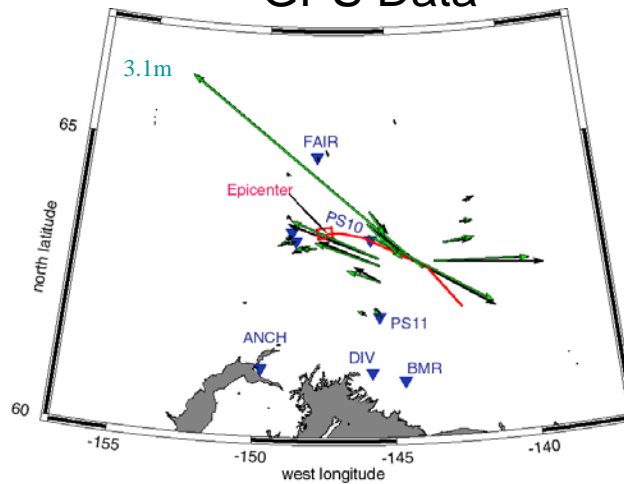
## Kinematic Model Cont.



## Construction of Kinematic Model for Denali Earthquake



## GPS Data

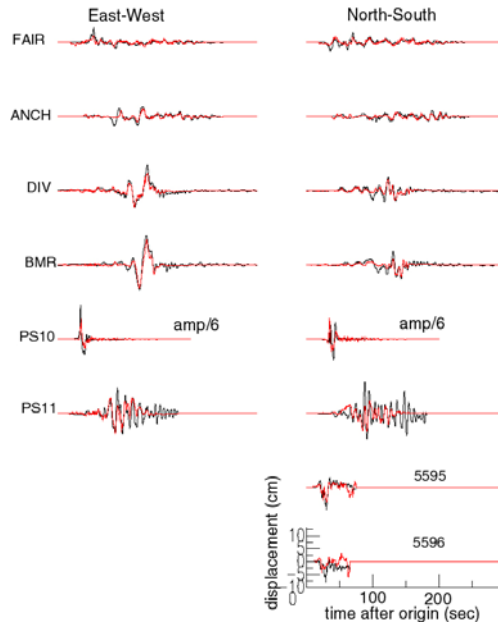


GPS data are from both the continuous Alaska Deformation Array, and campaign-mode observations of Hreinsdóttir et al. (2003). Black arrows show observed values and Green shows the values predicted by the model. Inverted triangles show the locations of the regional seismic stations. The red lines show the 5-segment fault model.

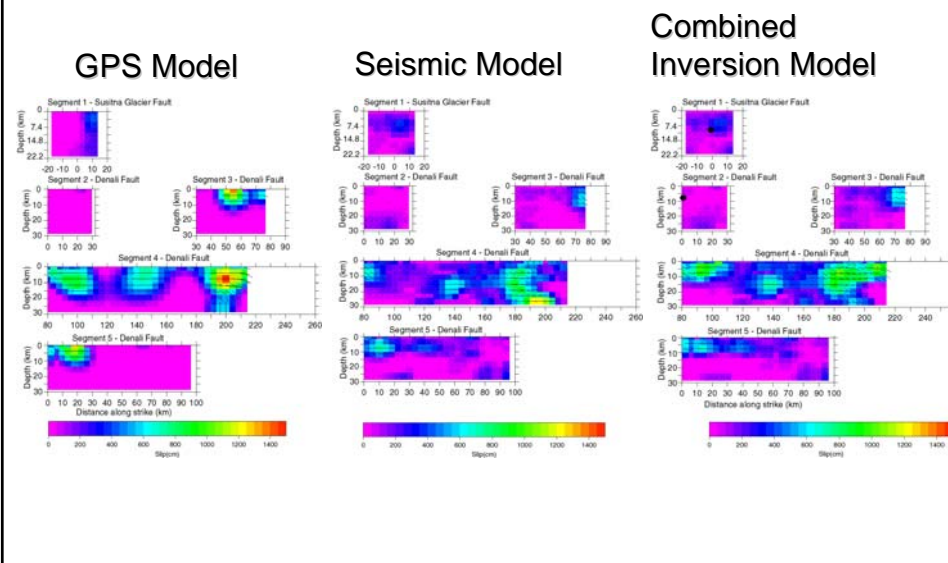
## Regional Seismic Waveform Data

Observed (black) and synthetic displacement waveforms (red) are compared. Station PS10 was not used to invert for the slip model and the fit is a forward prediction.

- Data Sources
  - US Geological Survey
  - Geophysical Institute, UA, Fairbanks
  - Alyeska
- Data from 8 stations were integrated to displacement, bandpass filtered between 0.01-0.5 Hz and resampled to 2 sps.

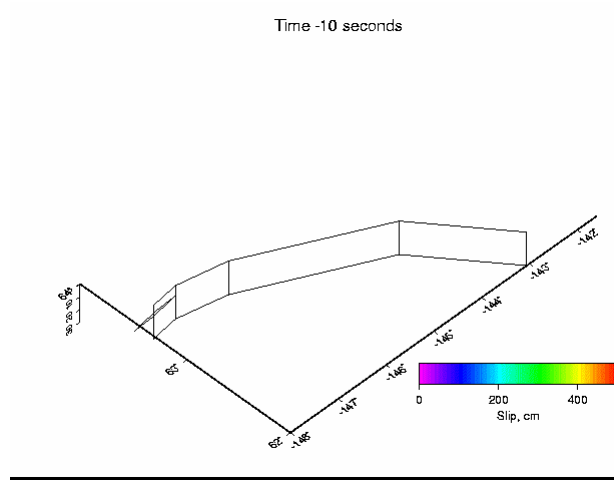


## Kinematic Inversion Results

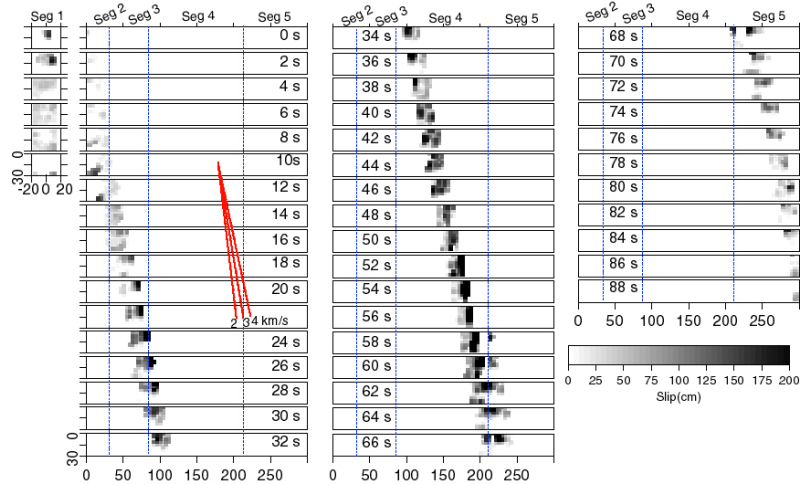




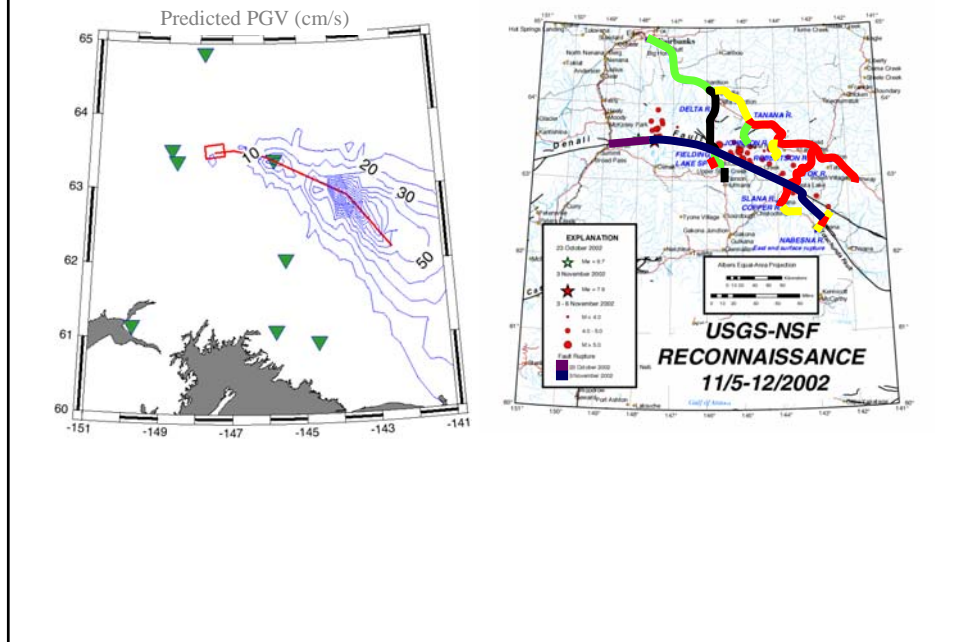
# Rupture Kinematics



# Rupture Kinematics



## Estimated PGV vs. Observed Liquefaction



## Realtime Finite-Source Applications