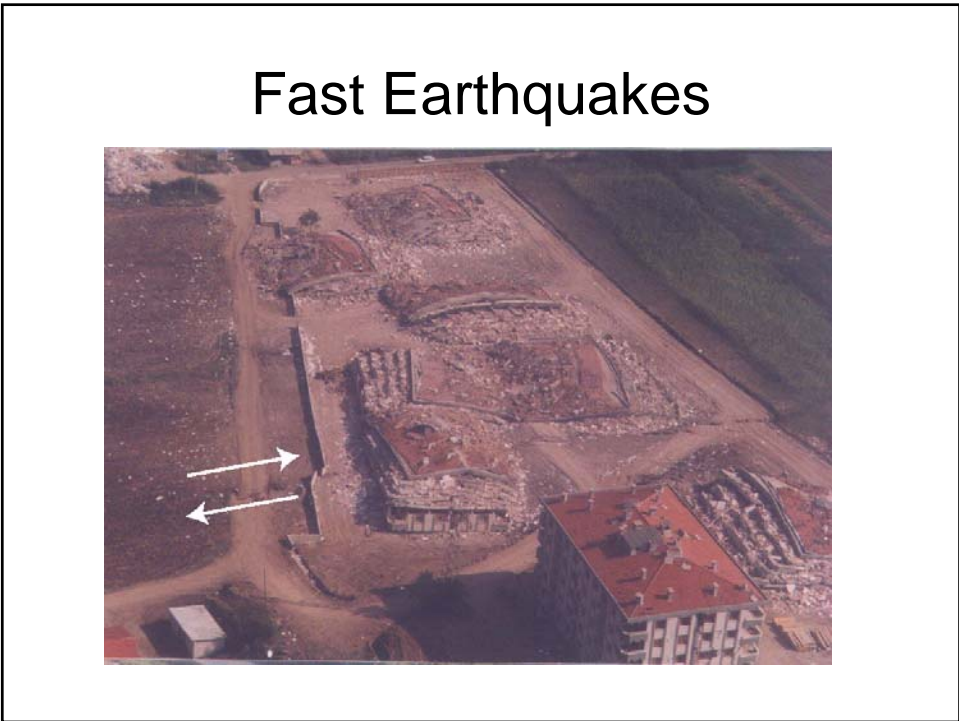
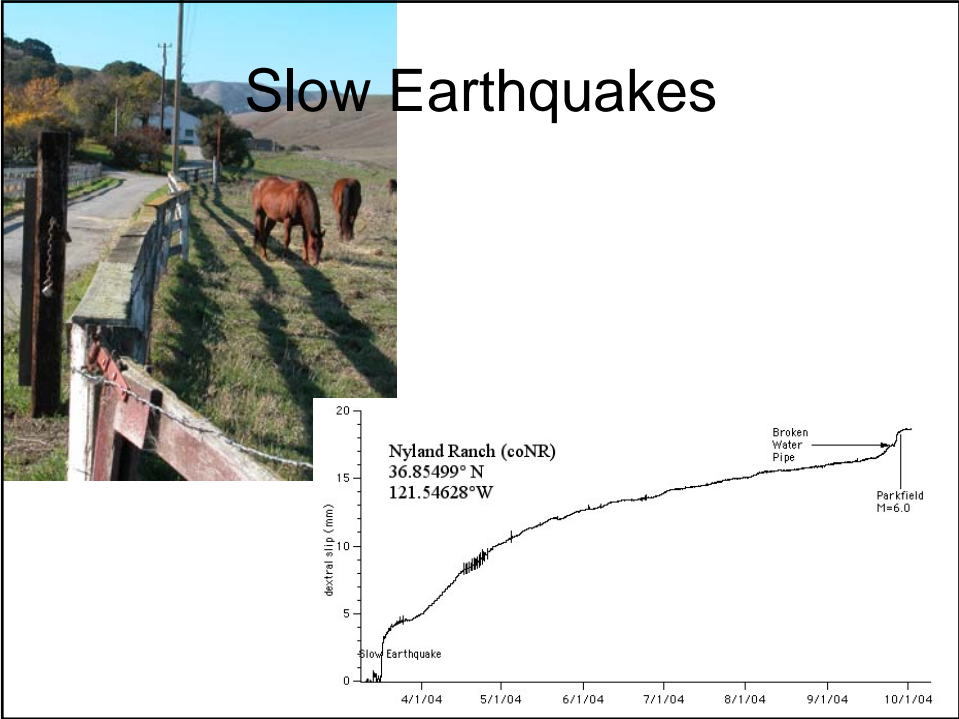


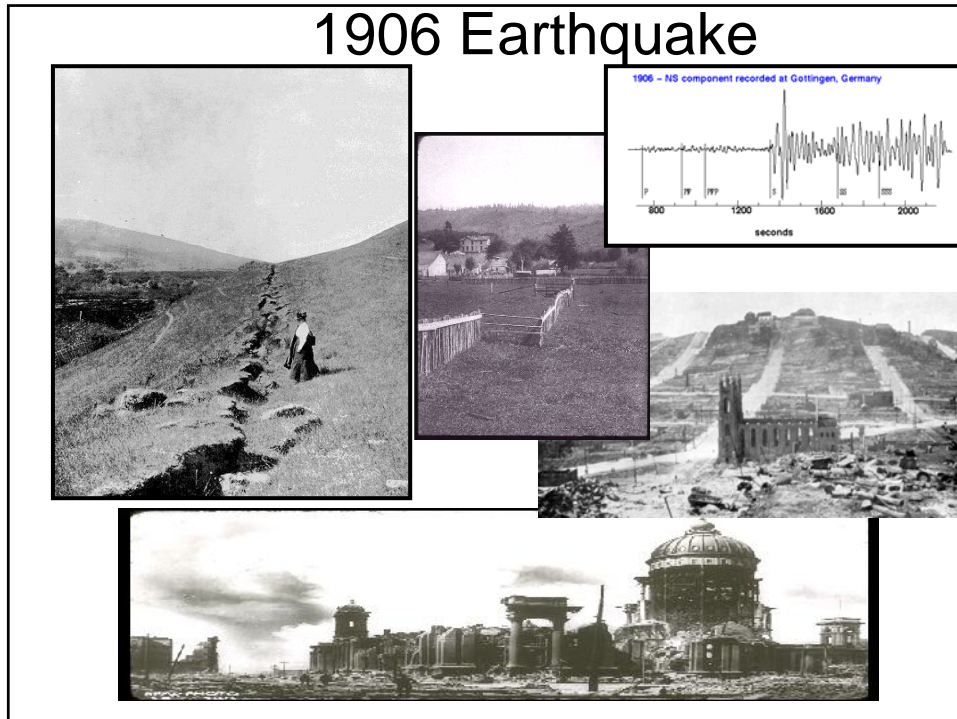
Overview of Seismotectonics

The relationship between
earthquake occurrence and
tectonic processes

- What is an earthquake?
 - Earth faulting event?
 - Propagating elastic waves that are felt and can cause damage?
- Where do earthquakes occur?
- Earthquake model & useful definitions
 - Elastic rebound theory – earthquake cycle
 - Scalar seismic moment
 - Moment magnitude
 - Faulting types and how to read focal mechanism diagrams
- Overview of plate tectonics
 - Evidence
 - Absolute and relative plate motions
 - Types of plate boundaries
 - Triple junctions
- Tectonic environments of some notable earthquakes



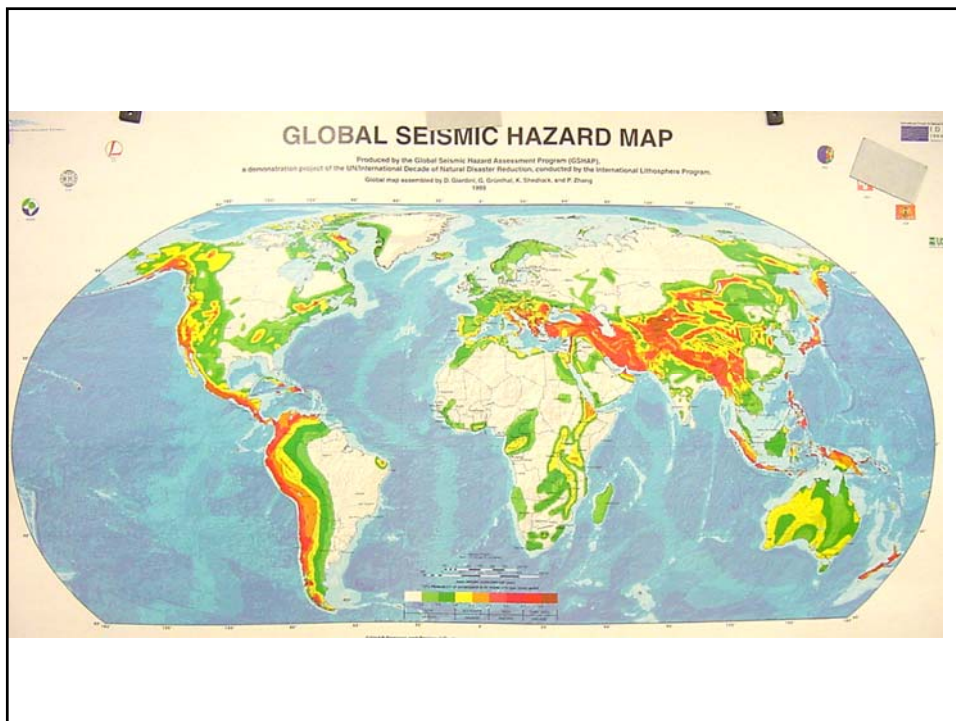
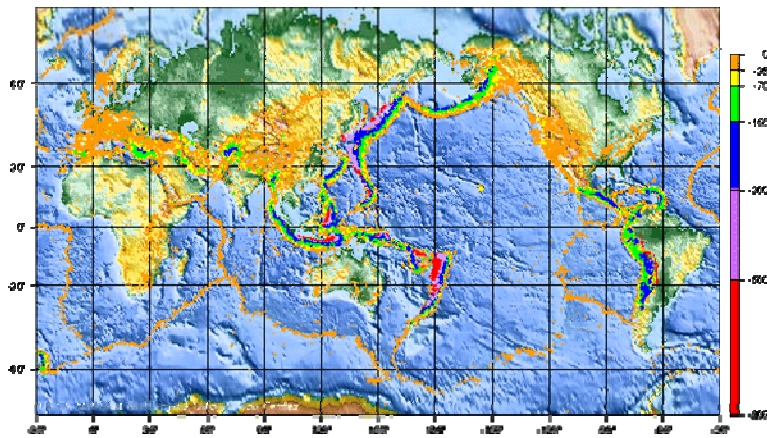
1906 Earthquake



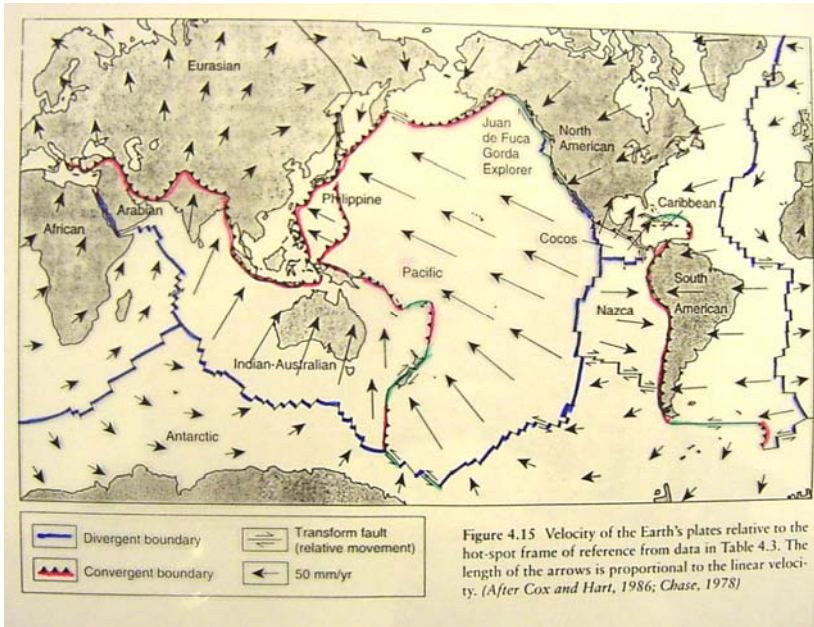
- What is an earthquake?
 - Earth faulting event?
 - Propagating elastic waves that are felt and can cause damage?
- Where do earthquakes occur?
- Earthquake model & useful definitions
 - Elastic rebound theory – earthquake cycle
 - Scalar seismic moment
 - Moment magnitude
 - Faulting types and how to read focal mechanism diagrams
- Overview of plate tectonics
 - Evidence
 - Absolute and relative plate motions
 - Types of plate boundaries
 - Triple junctions
- Tectonic environments of some notable earthquakes

Where do earthquakes occur?

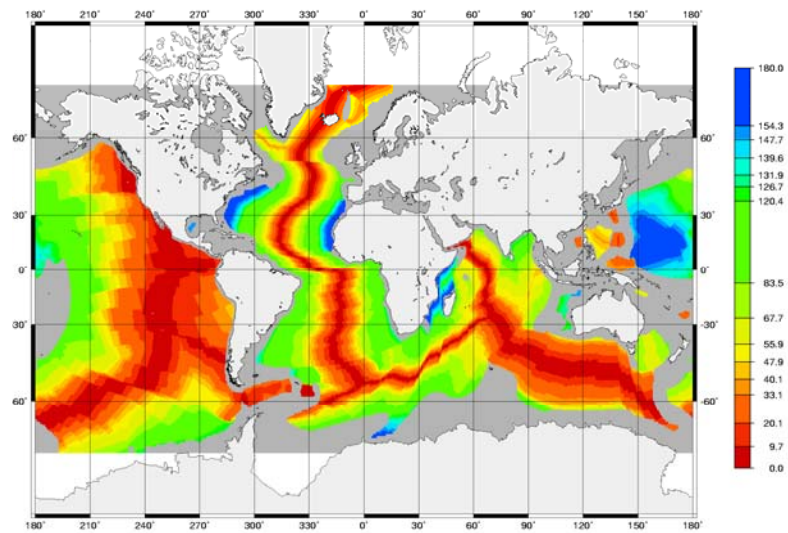
World Seismicity 1990 - 2000



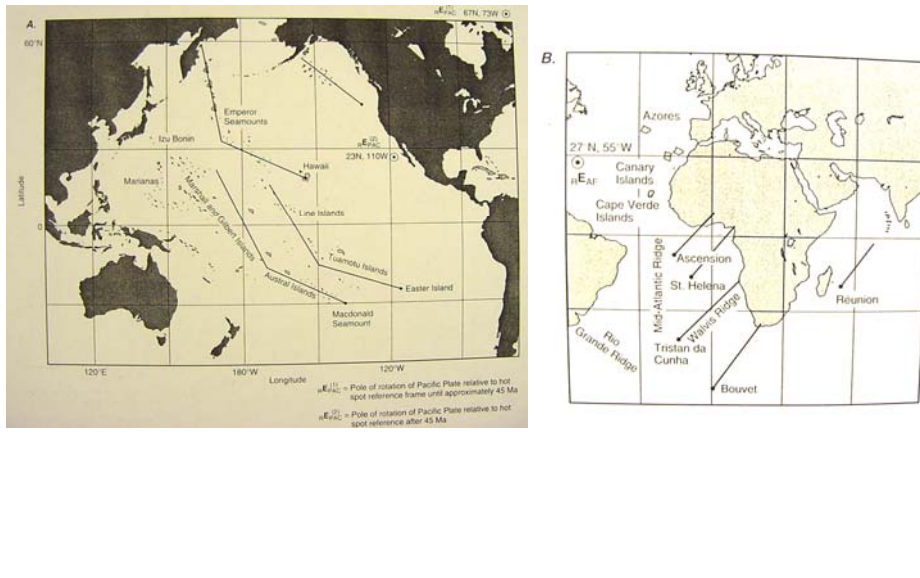
Absolute Plate Motion and Plate Boundary Type



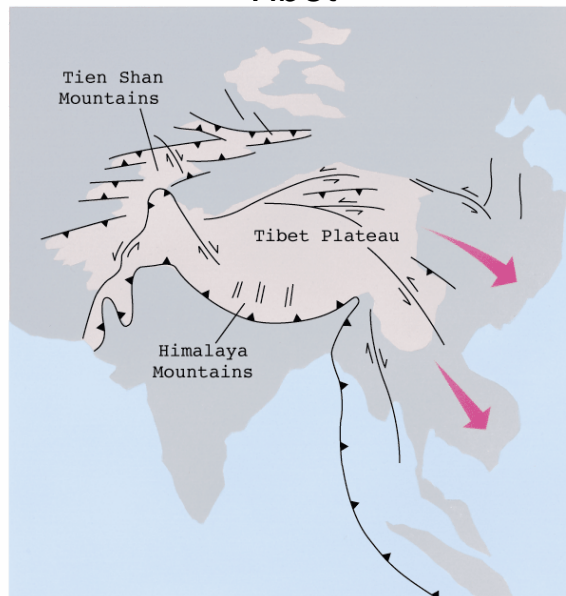
Age of the ocean floor & magnetic stripes



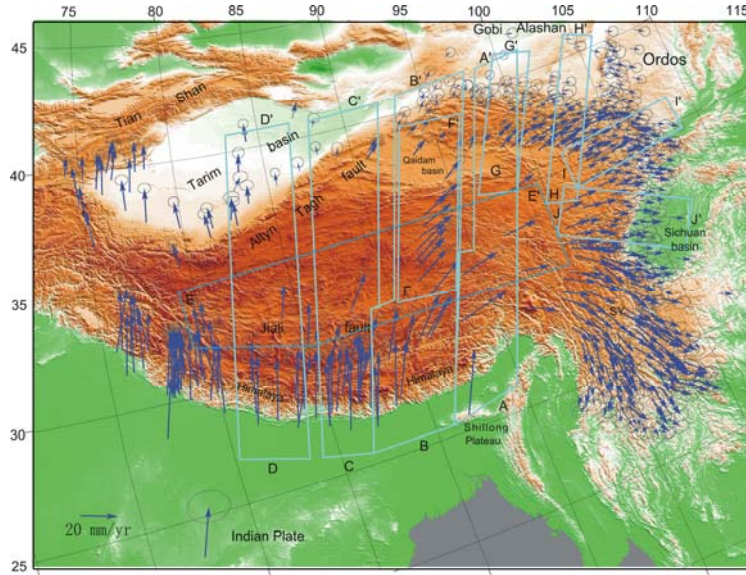
Hot Spot Tracks



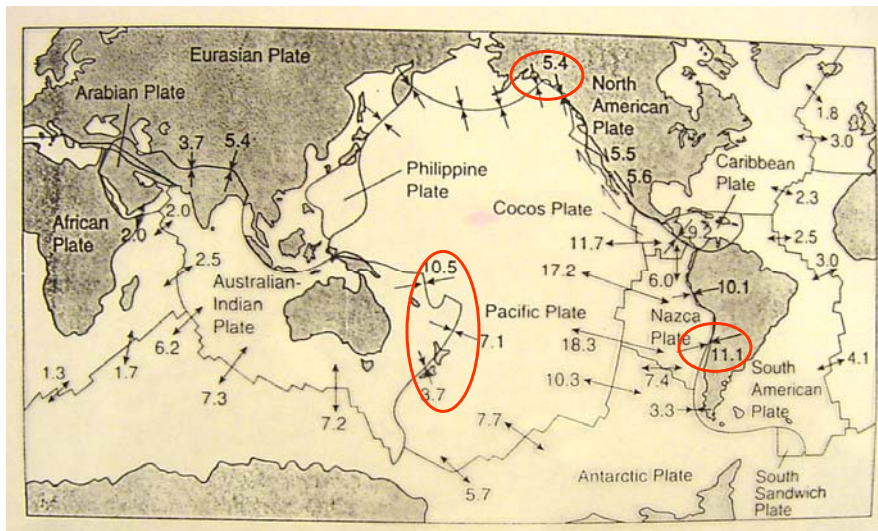
GPS Deformation – Escape Tectonics of Tibet



GPS Deformation – Escape Tectonics of Tibet

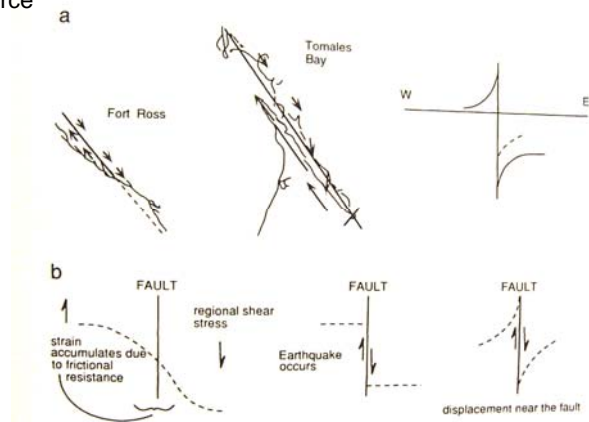


Relative Plate Motions (cm/yr)



- What is an earthquake?
 - Earth faulting event?
 - Propagating elastic waves that are felt and can cause damage?
- Where do earthquakes occur?
- Earthquake model & useful definitions
 - Elastic rebound theory – earthquake cycle
 - Scalar seismic moment
 - Moment magnitude
 - Faulting types and how to read focal mechanism diagrams
- Overview of plate tectonics
 - Evidence
 - Absolute and relative plate motions
 - Types of plate boundaries
 - Triple junctions
- Tectonic environments of some notable earthquakes

- Earthquake model
 - Elastic rebound theory (H. F. Reid, 1910)
 - Necessary conditions
 - Deformable elastic medium
 - Brittle rheology, or
 - Frictional resistance on pre-existing fault
 - Outside driving force



The Earthquake Cycle

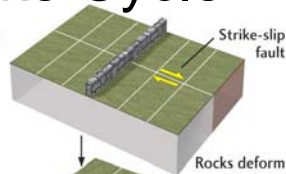
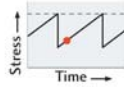
ROCKS DEFORM ELASTICALLY, THEN REBOUND DURING AN EARTHQUAKE RUPTURE

Steady elastic strain accumulation

Elastic rebound

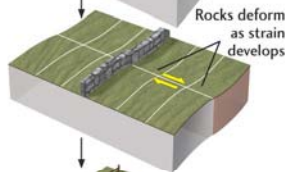
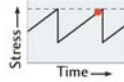
TIME 1

A farmer builds a stone wall across a right-lateral strike-slip fault a few years after its last rupture.



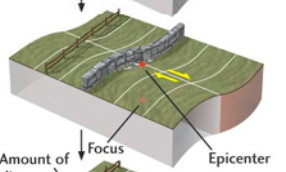
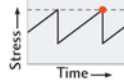
TIME 2

Over the next 150 years, the relative motion between blocks on either side of the locked fault causes the ground and the stone wall to deform.



TIME 3

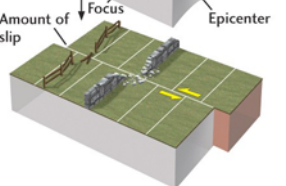
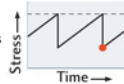
Just before the next rupture, a new fence is built across the already-deformed land.



When the stress exceeds the strength of the fault, a rupture begins at the first point of failure—the focus—beneath the epicenter on the surface. The rupture expands rapidly across the fault.

TIME 4

The rupture displaces the fault, lowering the stress, and the elastic rebound restores the blocks to their pre-stressed state. Both the rock wall and the fence are shifted equal amounts across the fault trace. The rebound straightens the rock wall, but the fence exhibits a reverse curve.



The Earthquake Cycle



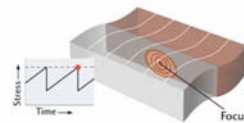
Elastic rebound

When the stress exceeds the strength of the fault, a rupture begins at the first point of failure—the focus—beneath the epicenter on the surface. The rupture expands rapidly across the fault.

TIME 4

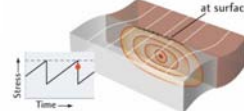
The rupture displaces the fault, lowering the stress, and the elastic rebound restores the blocks to their pre-stressed state. Both the rock wall and the fence are shifted equal amounts across the fault trace. The rebound straightens the rock wall, but the fence exhibits a reverse curve.

0 Second
Rupture expands circularly on fault plane, sending out seismic waves in all directions.

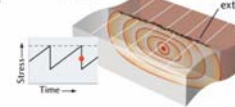


A fence built across the San Andreas fault near Bolinas, California, is offset by nearly 3 m after the great San Francisco earthquake

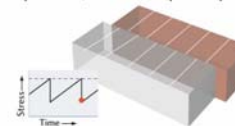
5 Seconds
Rupture continues to expand as a crack along the fault plane. When rupture front reaches the surface, displacements occur along the fault trace and rocks at the surface begin to rebound from their deformed state.



10 Seconds
Rupture front progresses down the fault plane, reducing the stress and allowing rocks on either side to rebound. Seismic waves continue to be emitted in all directions as the fault propagates.



20 Seconds
Rupture has progressed along the entire length of the fault. The fault has reached its maximum displacement, and the earthquake stops.



Friction: Stick-Slip vs. Creep

- Earthquakes in the Lab:

Earthquake like behavior

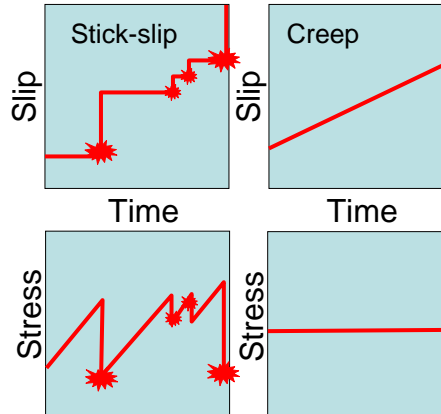
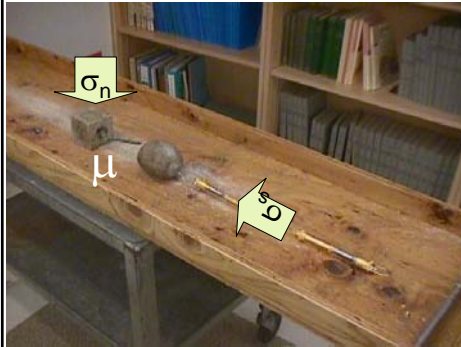
Strength of faults

- Increasing normal stress increases frictional strength

↳ Amonton's law $\sigma_s = \mu\sigma_n$

- Earthquakes indicate stick slip

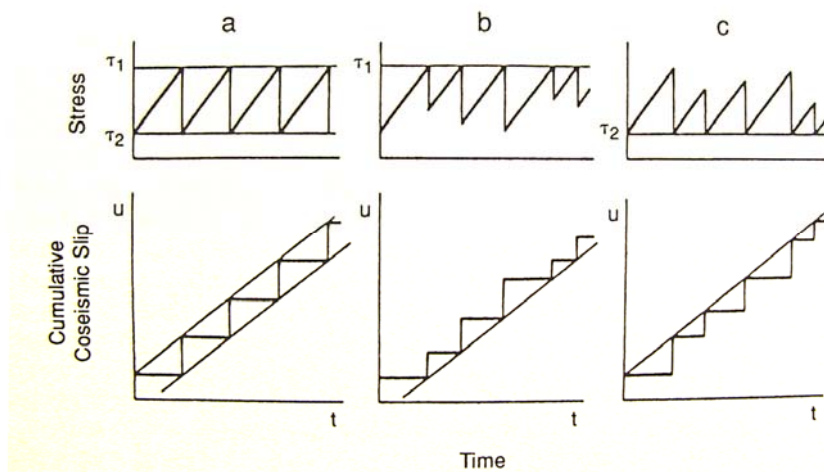
- Some faults (including the Hayward fault) also have steady creep



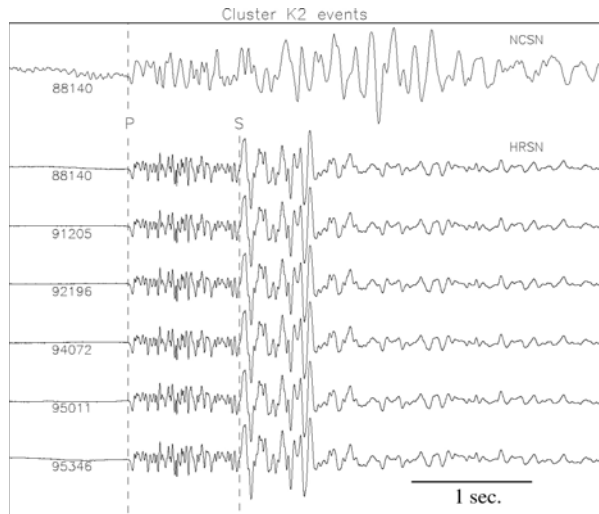
Characteristic/Periodic

Time Predictable

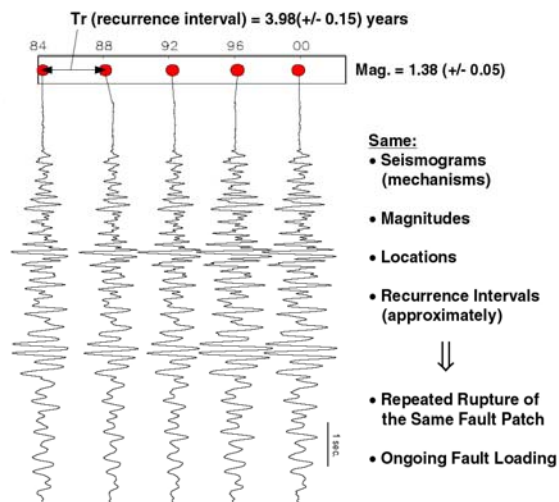
Slip Predictable



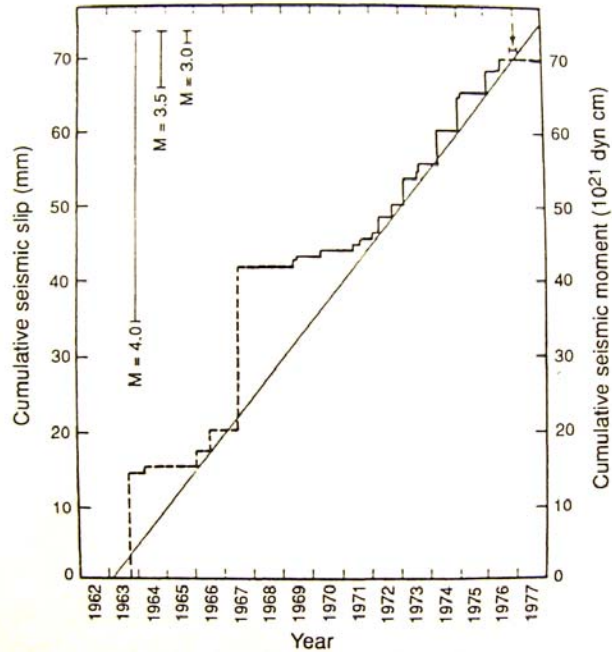
Repeating Micro Earthquakes Display Characteristic Earthquake Behavior



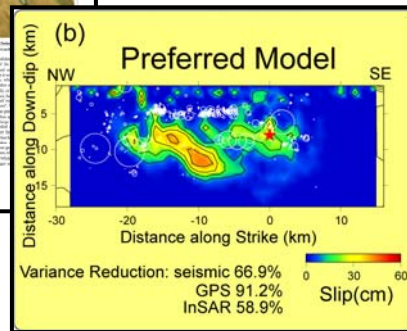
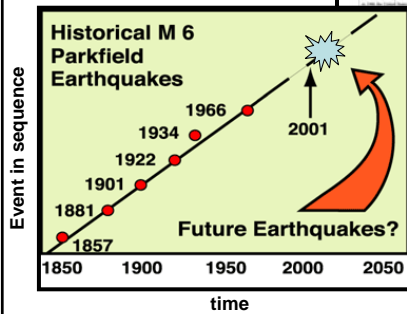
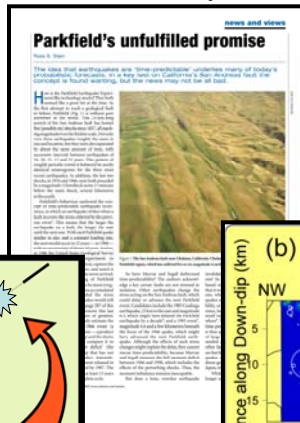
Characteristically Repeating Micro-quakes: (Example from the Hayward Fault.)



Cumulative slip on the Calaveras fault from 1962-1978 displays time predictable behavior



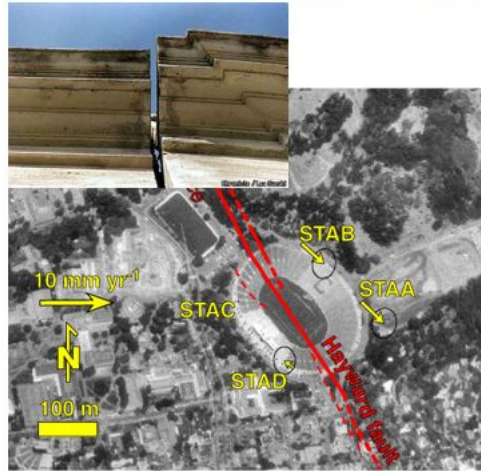
Unfortunately the 2004 Parkfield Earthquake Displayed Neither Time or Slip Predictable Behavior



Earthquakes vs. Fault Creep

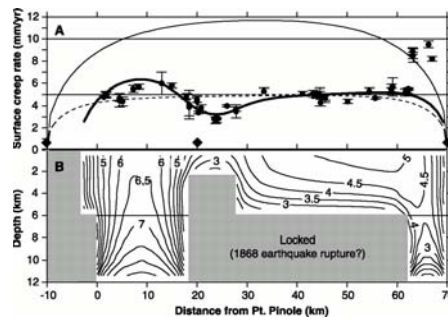
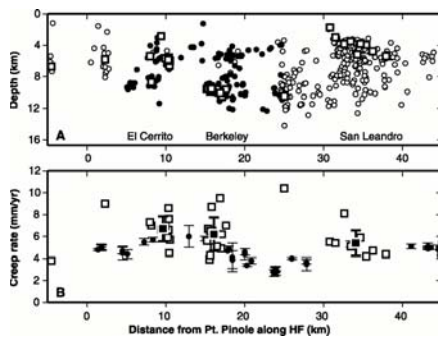


Hayward, 1868 and 1971



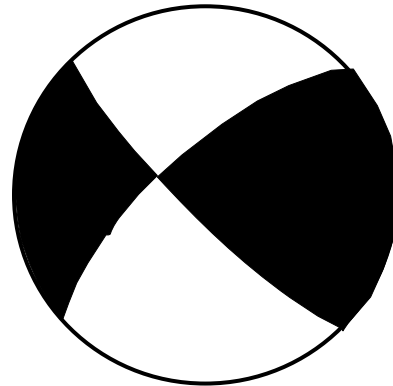
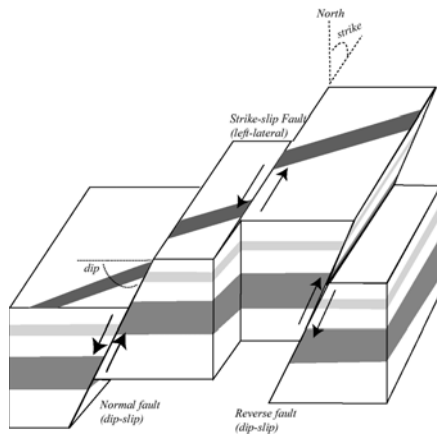
www.mcs.csuhayward.edu/~shirschf/tour-4.html

Hayward Fault Repeating Earthquakes and Slip Rate



Burgmann et al. 2000 - estimates of Hayward fault creep rate from geodetic and characteristically repeating micro earthquakes suggests the seismic potential of the NHF may be less than previously thought.

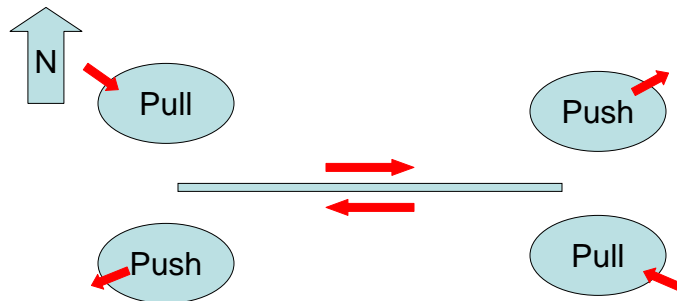
Types of Faulting



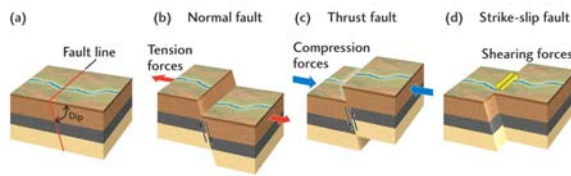
How does this relate to

this?

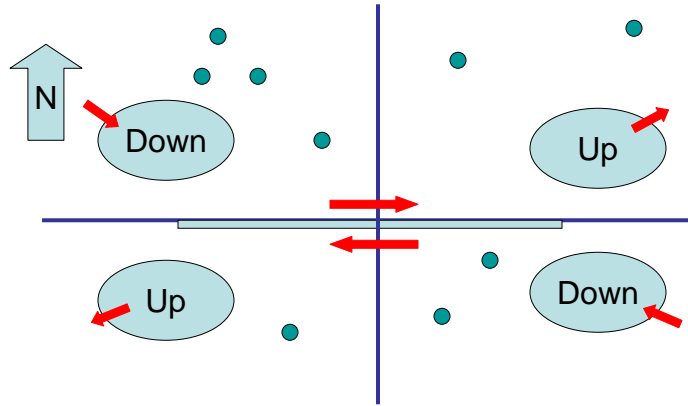
Sign of First Motion Earthquake Geometry



Fault motion of different kinds of faults (normal, reverse, strike-slip) will produce distinctive seismic wave characteristics

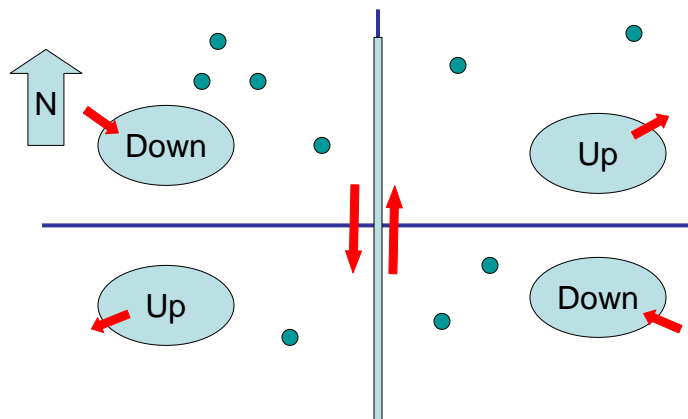


Earthquake Geometry



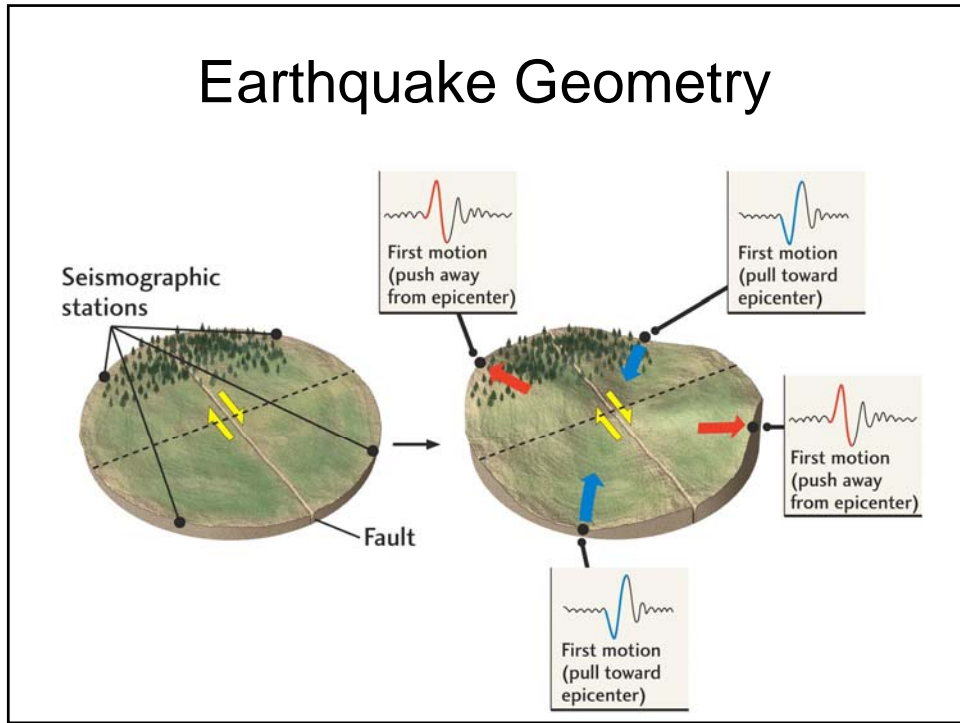
- First motion on vertical seismograph

Earthquake Geometry

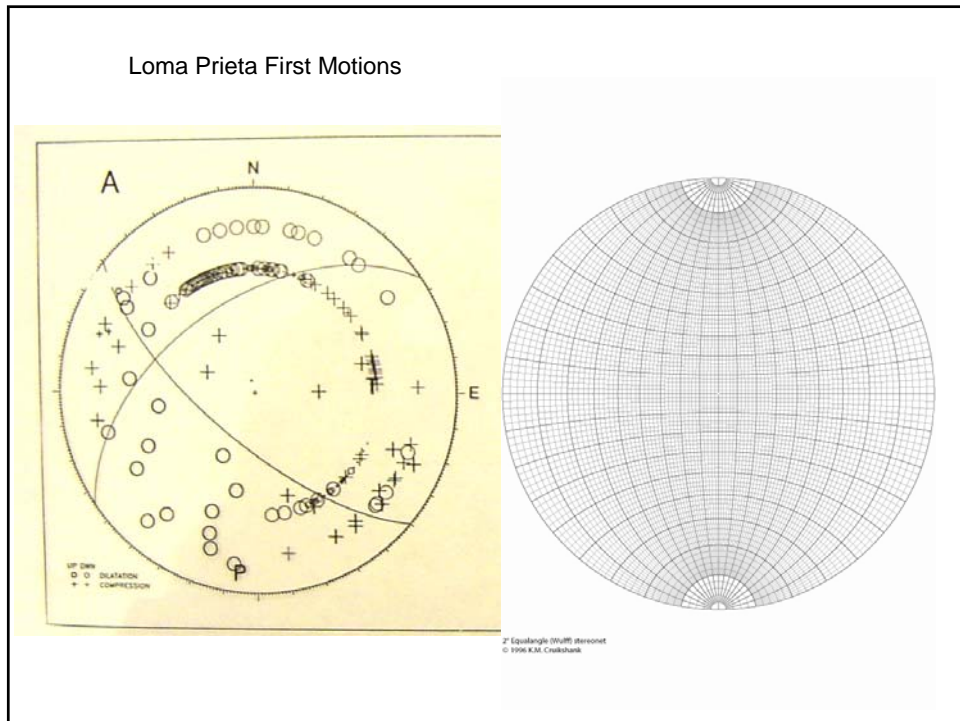


- First motion on vertical seismograph

Earthquake Geometry

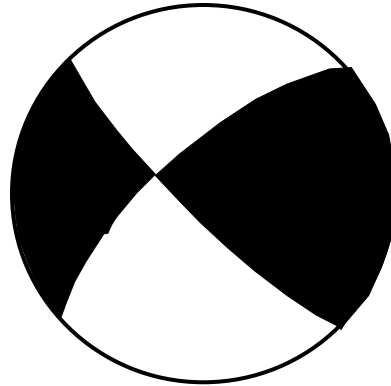


Loma Prieta First Motions



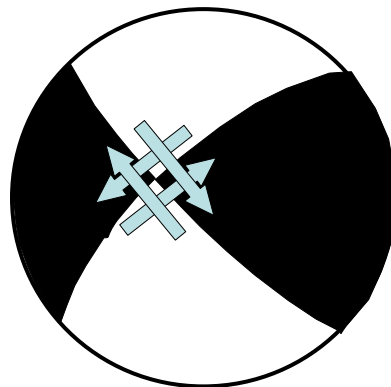
Earthquake Geometry – The Focal Mechanism

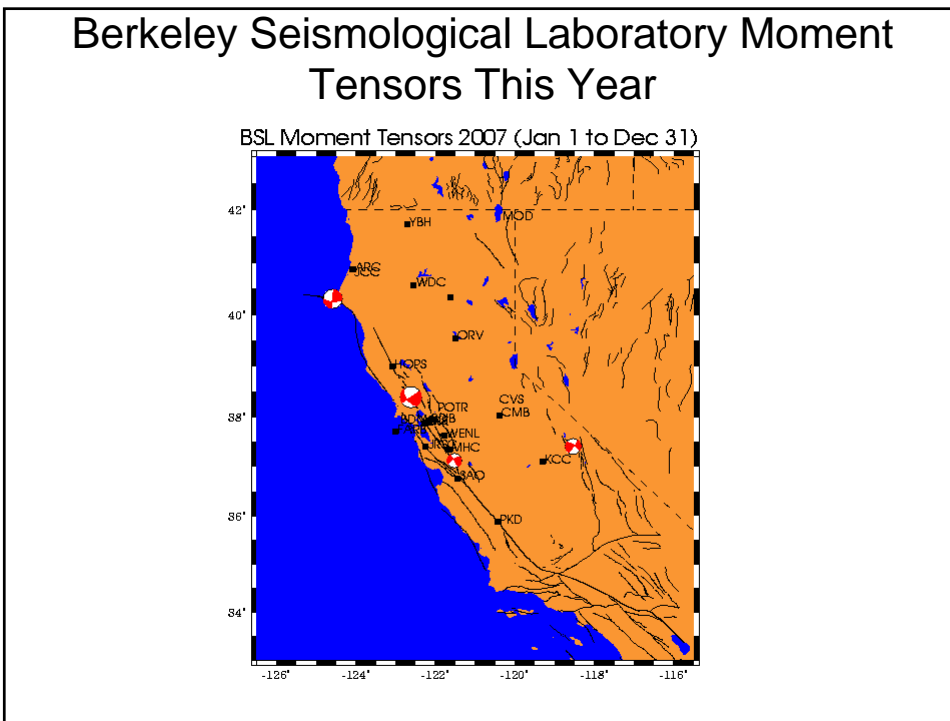
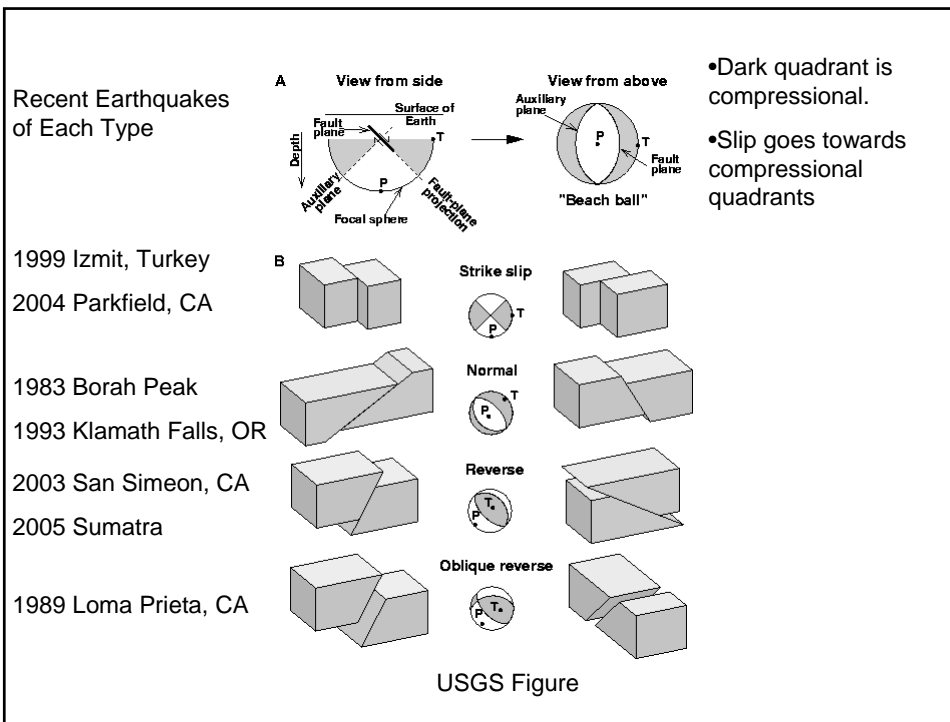
- The second nodal plane is at 90° from the first
- Draw the **Up** quadrant sectors solid



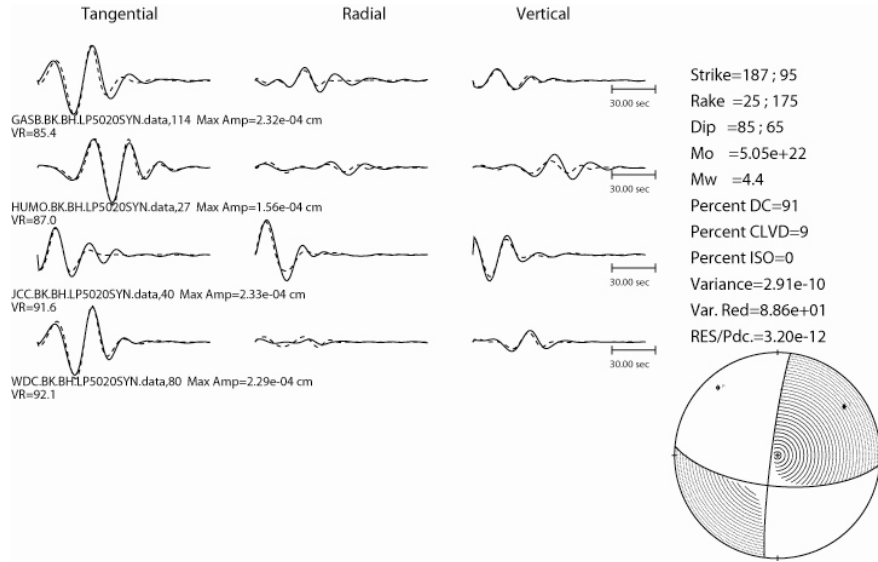
Earthquake Geometry – The Focal Mechanism

- The second nodal plane is at 90° from the first
- Draw the **Up** quadrant sectors solid
- A focal mechanism always indicates 2 possible fault planes
- How do we figure out which is the actual rupture?



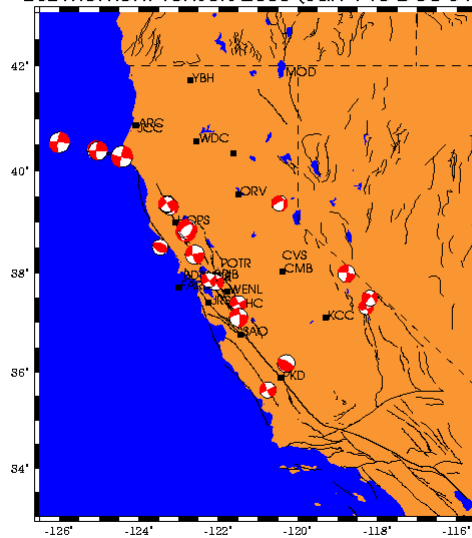


January 23, 2007 Mw4.4 Mendocino

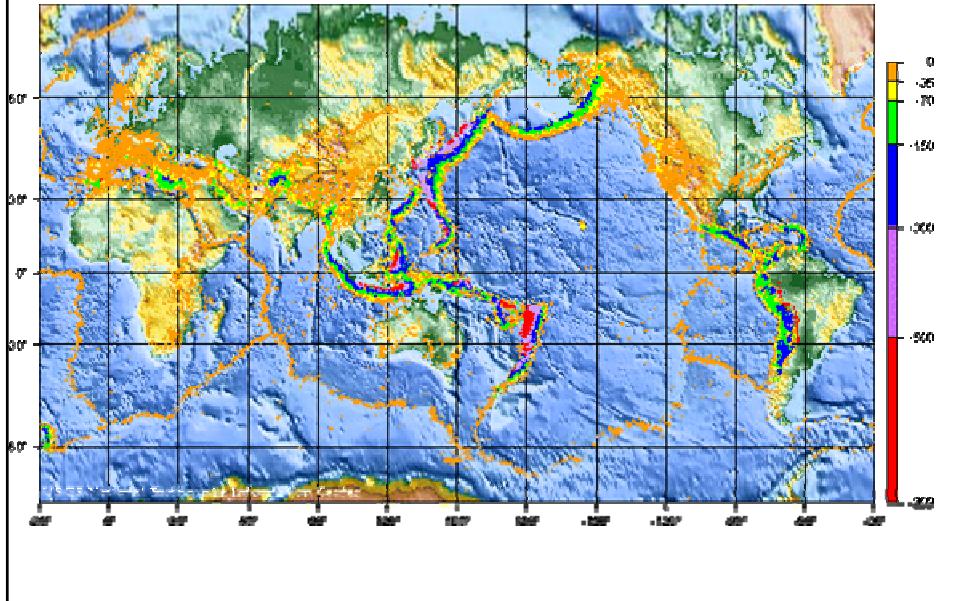


And last year

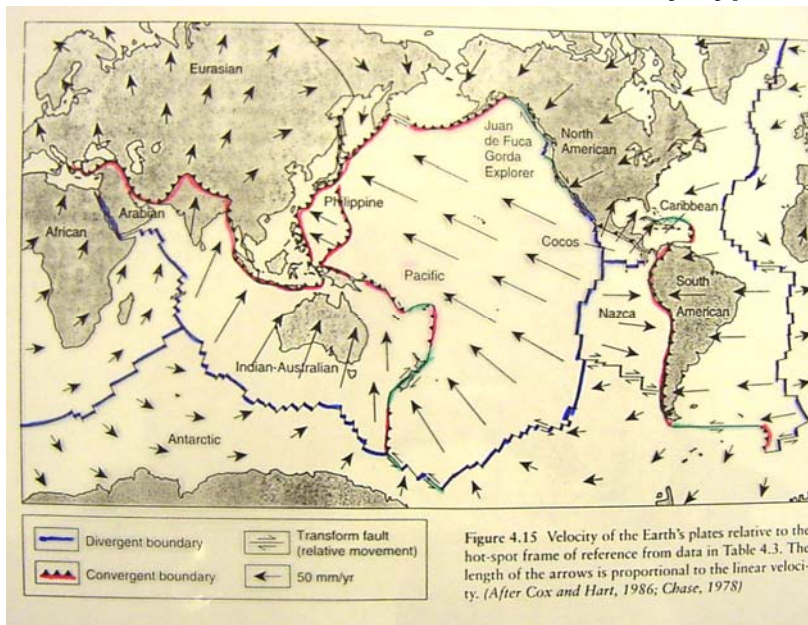
BSL Moment Tensors 2006 (Jan 1 to Dec 31)

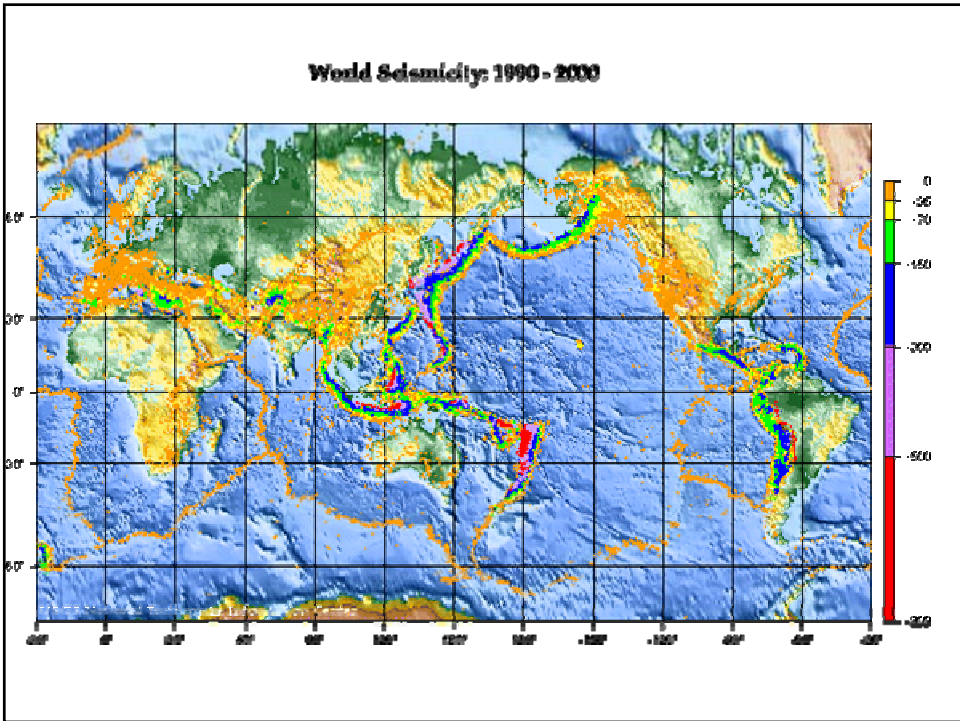
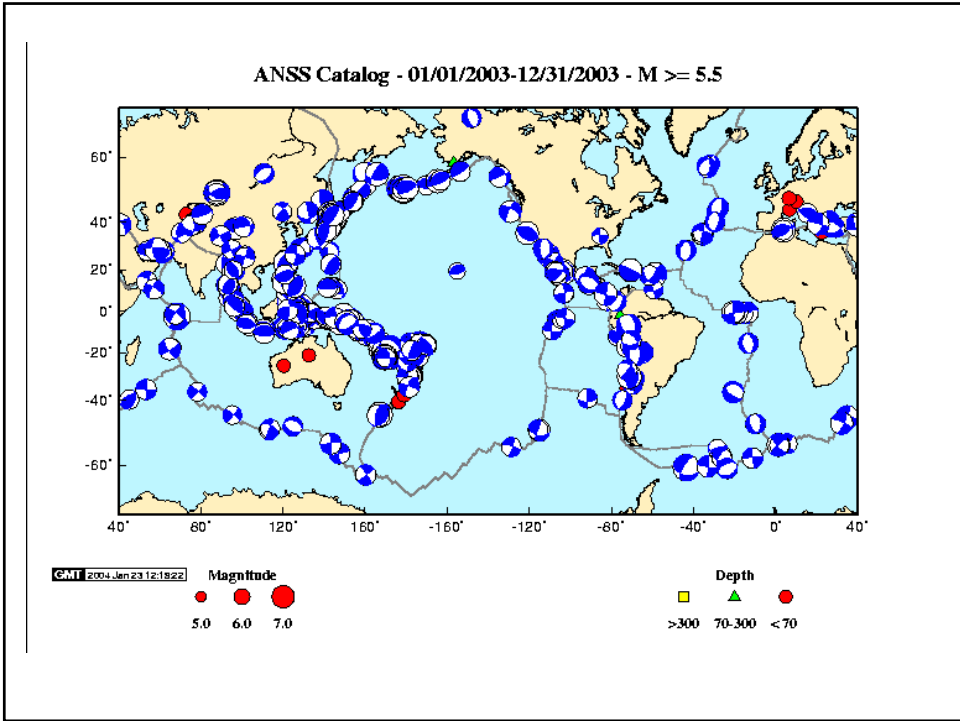


World Seismicity 1990 - 2000



Absolute Plate Motion and Plate Boundary Type





Basic Quantification

- Seismic Potency
 - Potency=area*slip
- Scalar Seismic Moment
 - $M_0 = \text{rigidity} \cdot \text{area} \cdot \text{slip} = \mu AD$
 - Has units of energy and is proportional to energy through the stress drop (change) h
- Moment Rate
 - $\dot{M}_0 = \mu A \dot{D}$
- Magnitude
 - $M_w = 2/3 \cdot \log(M_0) - 10.7$ or
 - $\log(M_0) = 1.5M_w + 16.05$ (a unit of magnitude increase corresponds to a 31-fold increase in energy)
- Statistics
 - Gutenberg-Richter
 - Omori Law

Size - Frequency Relationship

Frequency of Occurrence of Earthquakes Based on Observations since 1900

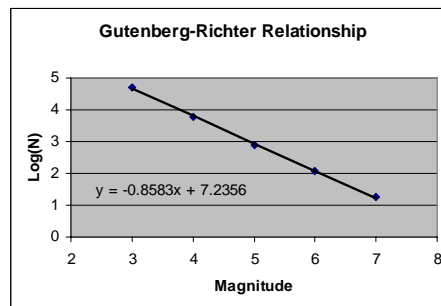
Great	> 8		higher	1/year
Major	7	-	7.9	18
Strong	6	-	6.9	120
Moderate	5	-	5.9	800
Light	4	-	4.9	~6,200
Minor	3	-	3.9	~49,000
Micro	1	-	3	~9000/day

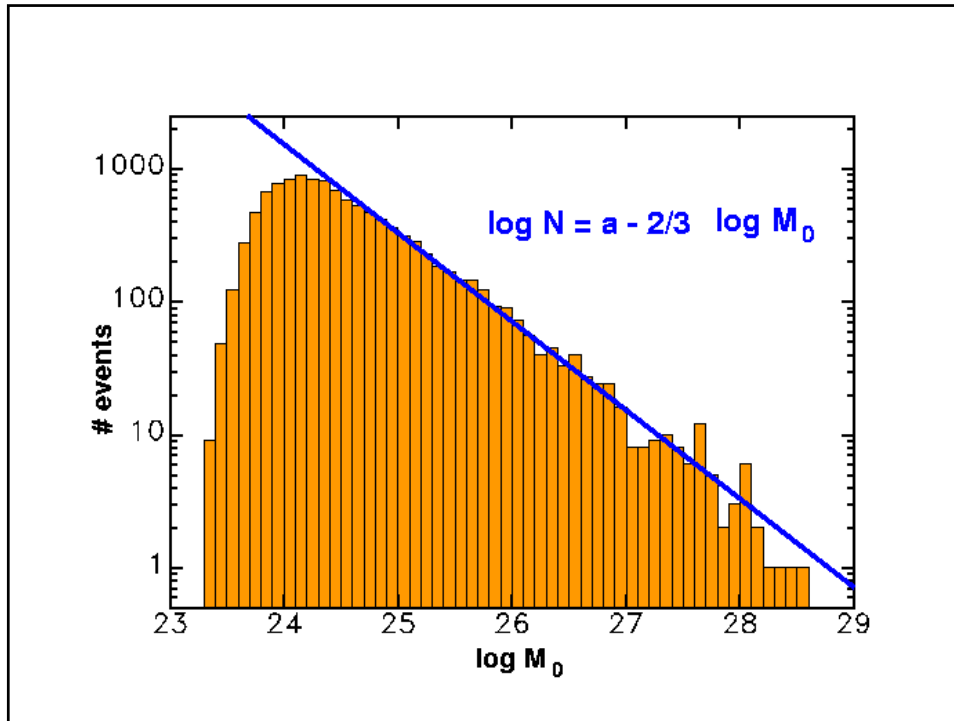
From neic.usgs.gov/neis/bulletin/mag7.htm#1999

Gutenberg-Richter Relationship:

$$\log(\text{number}) = a + b \cdot (\text{magnitude})$$

$$\log(N) = a + b \cdot M$$





Size - Frequency Relationship

Frequency of Occurrence of Earthquakes Based on Observations since 1900

Great	> 8	higher	1/year
Major	7	7.9	18
Strong	6	6.9	120
Moderate	5	5.9	800
Light	4	4.9	~6,200
Minor	3	3.9	~49,000
Micro	1	3	~9000/day

From neic.usgs.gov/neis/bulletin/mag7.html 1999

M7 ~ 1.99e+22 ergs

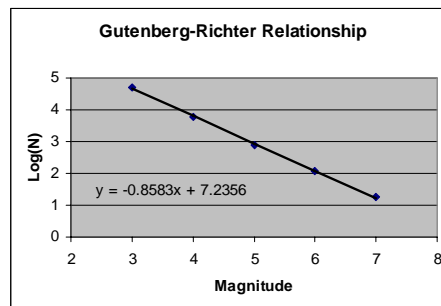
M6 ~ 6.30e+22 ergs ~ 1/30M7

It takes 32 M6 events to equal the moment of 1 M7, but M6 are only 10x more frequent

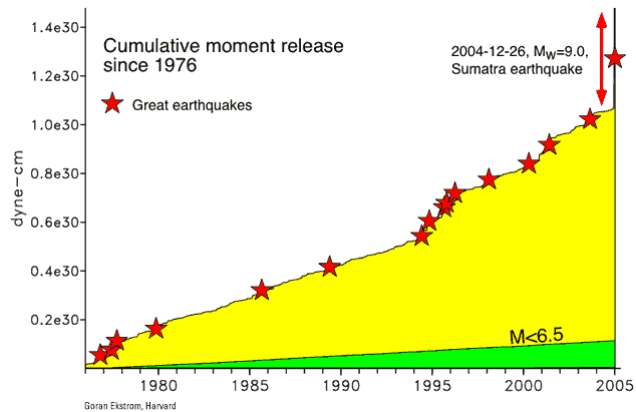
Gutenberg-Richter Relationship:

$\text{Log}(\text{number}) = a + b \cdot (\text{magnitude})$

$\text{Log}(N) = a + b \cdot M$

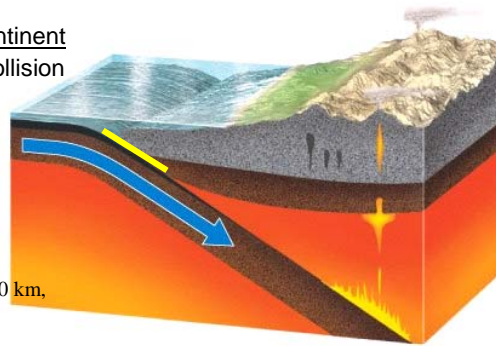


Cumulative Seismic Moment in Great Earthquakes



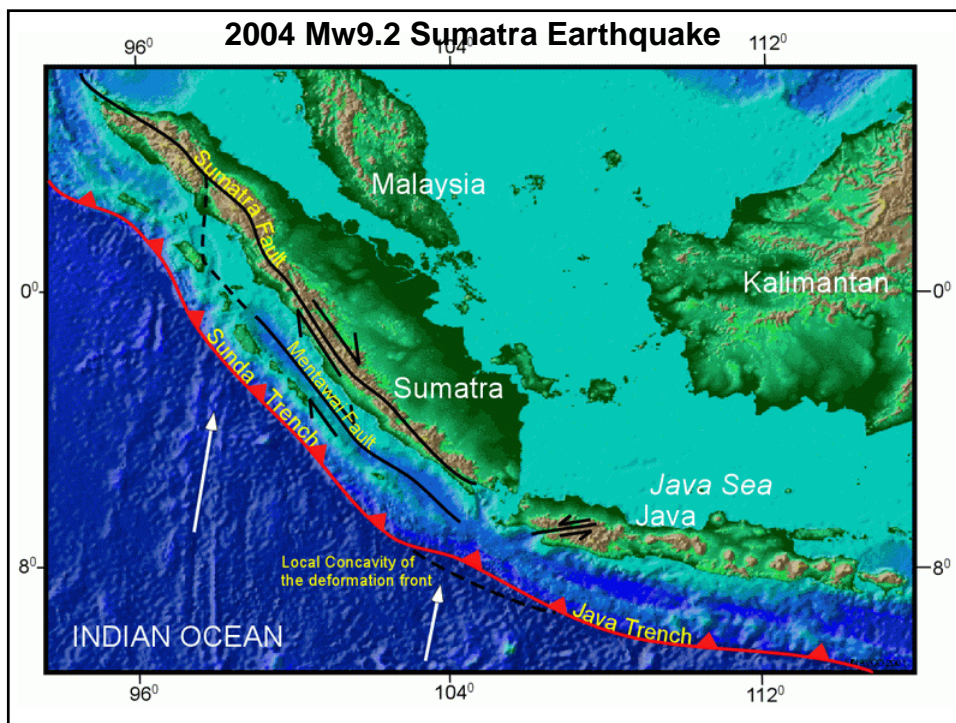
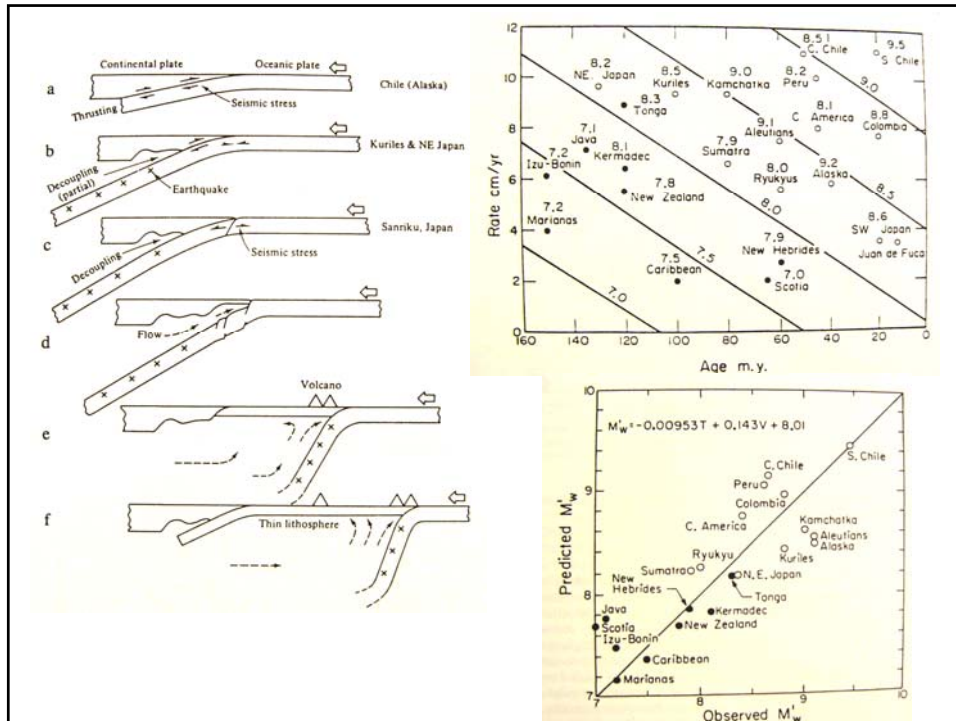
Convergent Boundaries

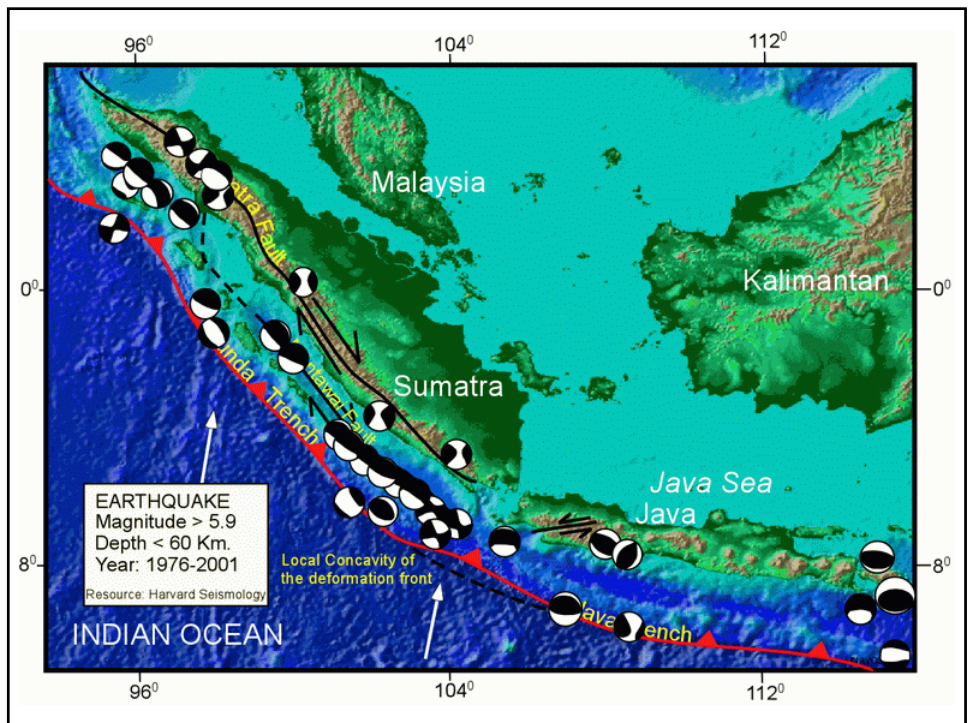
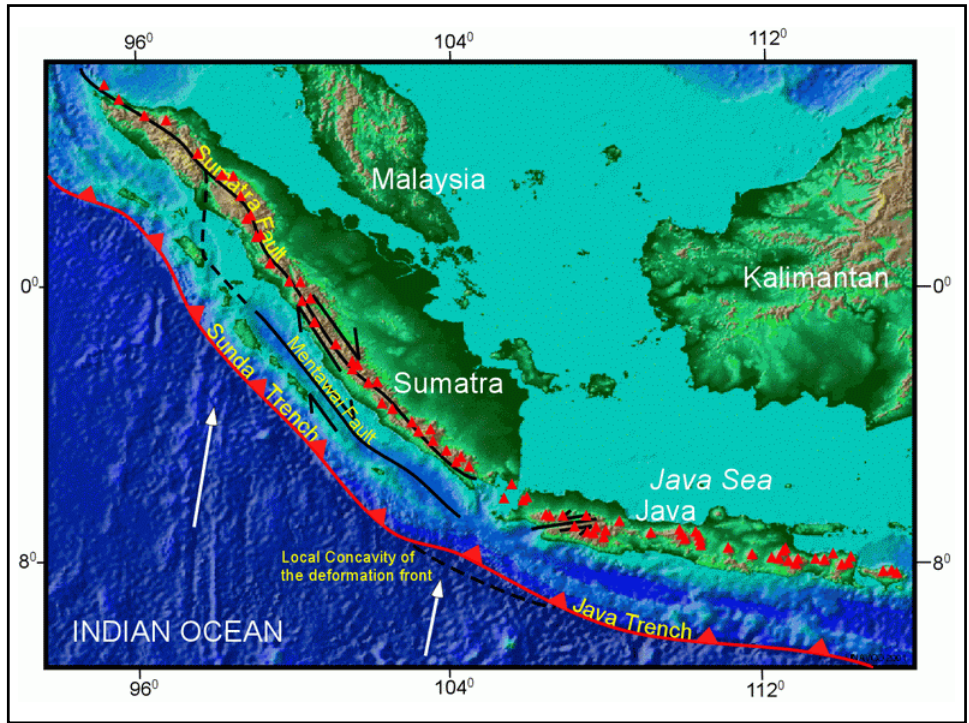
- Mostly ocean-ocean, and ocean-continent
- ~10% involve continent-continent collision
- Classic geometry

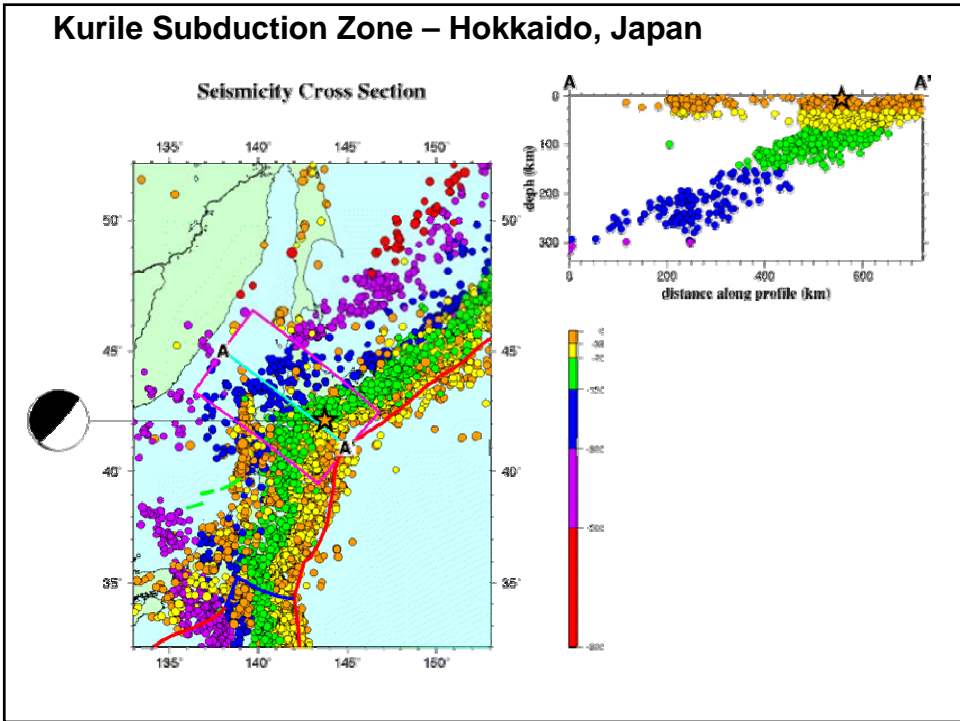
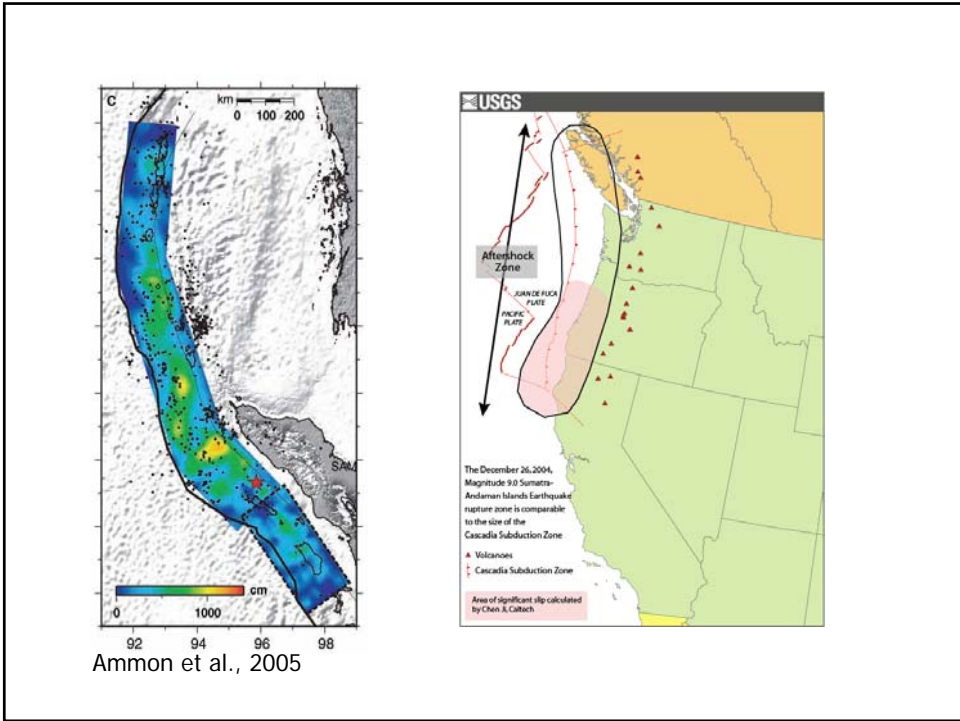


- Consider the Chilean EQ of 1960
 - $\mu=3.5e11$ dyne cm, $L=1000$ km, $W=200$ km, $D=22$ m
 - $M_0=1.5e30$ dyne cm and $M_w=9.4$
 - The moment rate of the same segment is $7.7e27$ dyne cm/year or a $M_w=7.9$ /year assuming a slip rate of 11 cm/yr
 - The recurrence time of $M_w9.4$ events is about 200 years

In subduction zones the shallow dipping fault (10-20 degrees) can lead to a large contact patch and therefore great earthquakes

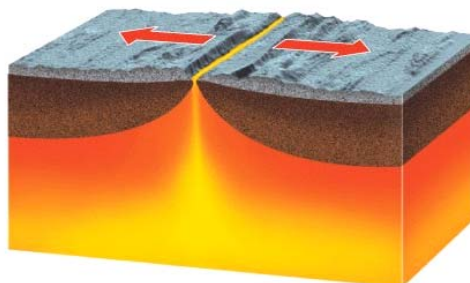




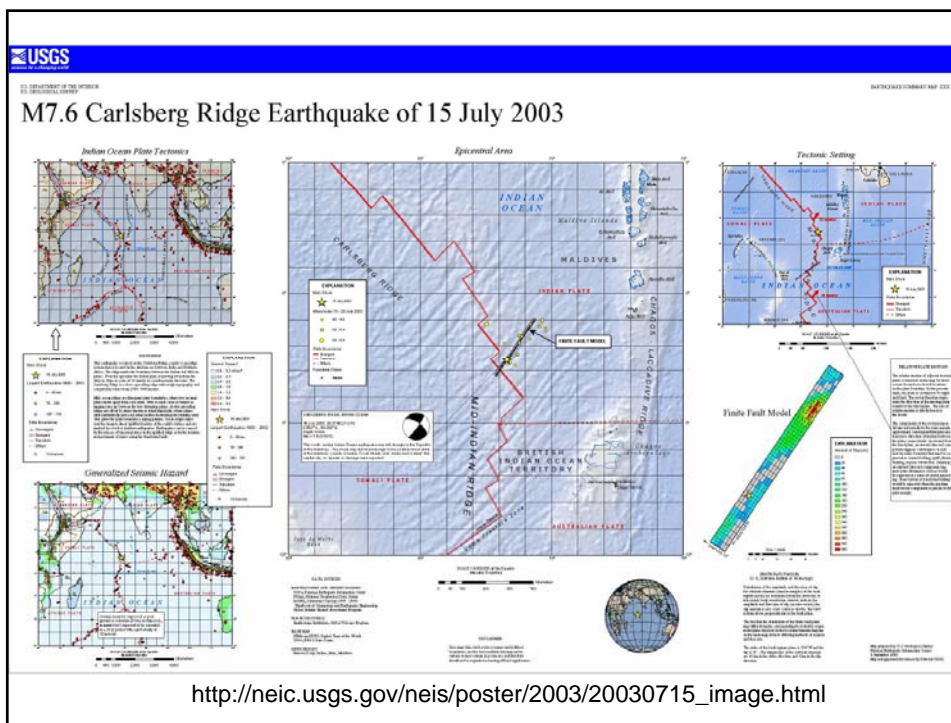


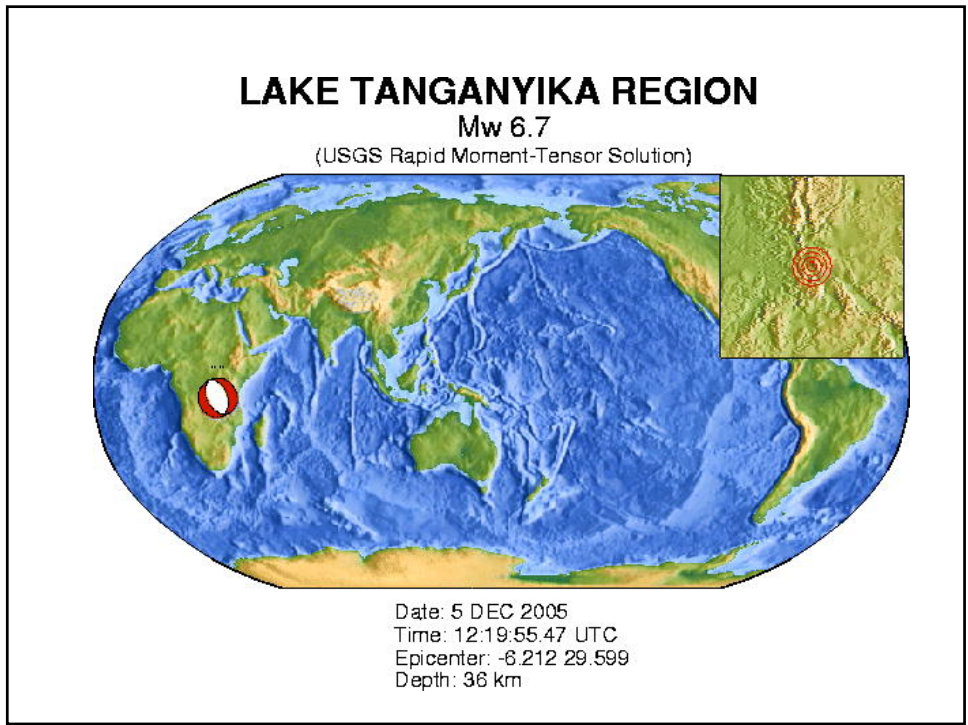
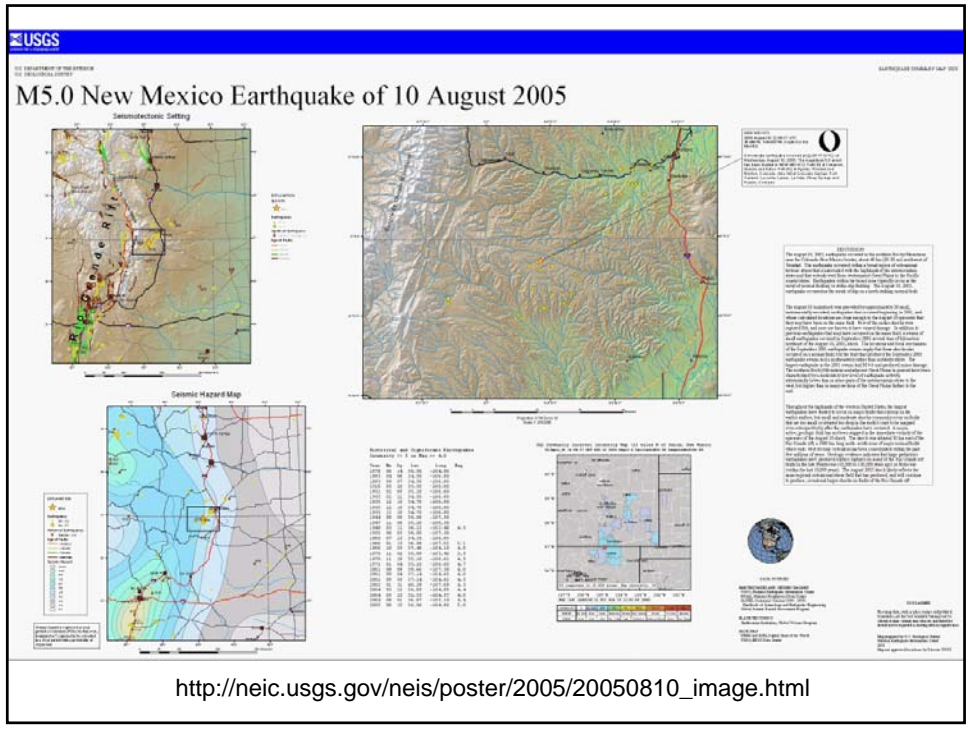
Divergent Boundaries

- Oceanic rifting
 - Normal-type events parallel to ridge axis
 - Strike-slip faulting on ridge-connenting transform faults
 - Normal events tend can be as large as 6-7 but tend to be small due to thin, warm crust
 - Transform events can be as large as magnitude 7-8
- Continental rifting
 - More complicated
 - Depth of faulting limited by relatively warm crust
 - Can get as large as magnitude 7 (e.g. 7.3 Borah Peak, Idaho)

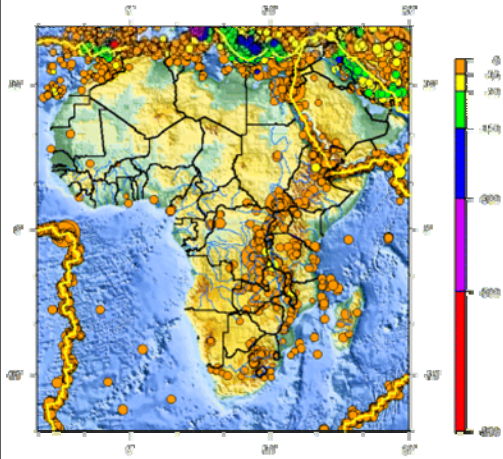


While fault length can grow normal events typical have dips close to 60-degrees and therefore the down-dip width is less than for reverse events

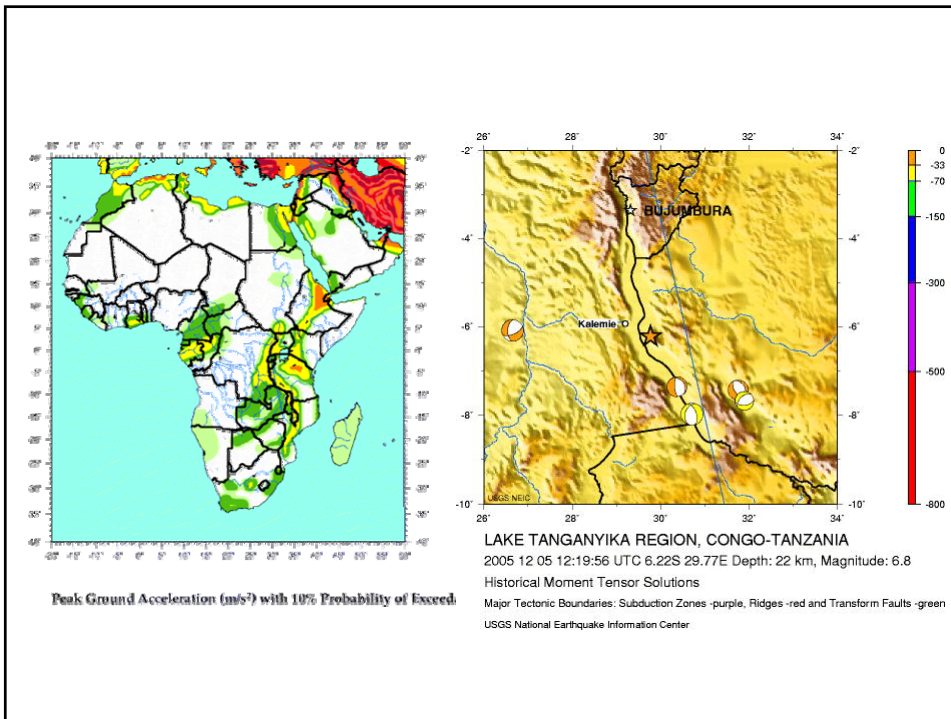
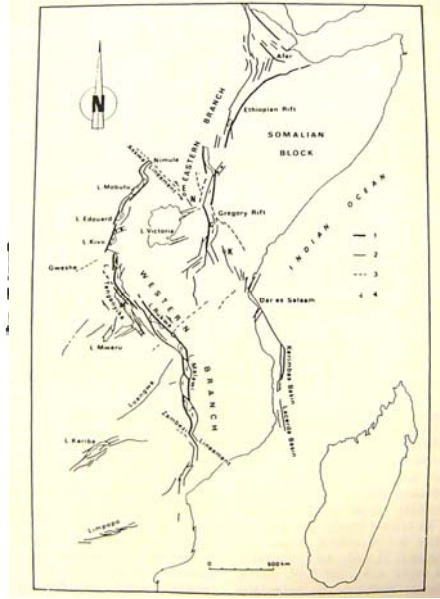




Red Sea and East African Rifting

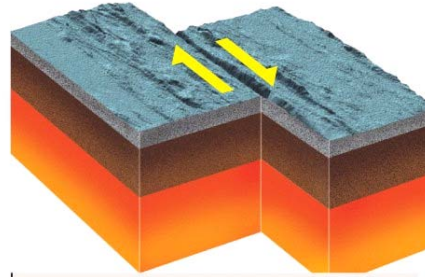


Seismicity of Africa, 1990 - 2000



Translational Boundaries

- Event size is limited by fault segmentation and depth of seismogenic zone
- SiO_2 begins to behave ductilely at 300°C corresponding to roughly 10 to 15 km depth
- 1906, $L=450\text{km}$, $W=15\text{km}$, $D=5\text{m}$ $M_0=1.2e28$ dyne cm and $M_w=8.0$
- Moment rate= $4.3e25$ dyne cm/yr or a $M_w=6.4/\text{yr}$ assuming a slip rate of 1.8 cm/yr
- Recurrence interval of ~ 280 yrs



Next Time

- Continue discussion of seismotectonics focusing on California
- Investigate how this information can be used in characterizing seismic hazard, earthquake forecasting and prediction.
- Study material in Chapter 11